



TWELFTH EDITION

Patrick L. Abbott

San Diego State University









NATURAL DISASTERS, TWELFTH EDITION

Published by McGraw Hill LLC, 1325 Avenue of the Americas, New York, NY 10019. Copyright © 2023 by McGraw Hill LLC. All rights reserved. Printed in the United States of America. Previous editions © 2020, 2017, and 2014. No part of this publication may be reproduced or distributed in any form or by any means, or stored in a database or retrieval system, without the prior written consent of McGraw Hill LLC, including, but not limited to, in any network or other electronic storage or transmission, or broadcast for distance learning.

Some ancillaries, including electronic and print components, may not be available to customers outside the United States

This book is printed on acid-free paper.

1 2 3 4 5 6 7 8 9 LWI 27 26 25 24 23 22

ISBN 978-1-264-09116-4 (bound edition) MHID 1-264-09116-8 (bound edition) ISBN 978-1-266-59223-2 (loose-leaf edition) MHID 1-266-59223-7 (loose-leaf edition)

Senior Director, Portfolio: Michelle Vogler Lead Product Developer: Lora Neyens Senior Marketing Manager: Kelly Brown Senior Content Project Manager: Vicki Krug Content Project Manager: Rachael Hillebrand

Senior Buyer: *Sandy Ludovissy* Lead Designer: *David W. Hash* Content Licensing Specialist: *Beth Cray*

 $Cover\ Image: \textit{NASA}\ Earth\ Observatory\ image\ by\ Joshua\ Stevens,\ using\ Landsat\ data\ from\ the\ U.S.$

 $Geological\ Survey,\ and\ MODIS\ data\ from\ NASA\ EOSDIS/LANCE\ and\ GIBS/Worldview$

Compositor: MPS Limited

All credits appearing on page or at the end of the book are considered to be an extension of the copyright page.

Library of Congress Cataloging-in-Publication Data

Names: Abbott, Patrick L., author.

Title: Natural disasters / Patrick L. Abbott.

Description: Twelfth edition. | New York : McGraw Hill LLC, 2023. |

Includes index.

Identifiers: LCCN 2021026407 | ISBN 9781264091164 (hardcover) | ISBN 9781266592232 (spiral bound) | ISBN 9781265125554 (hardcover)

Subjects: LCSH: Natural disasters.

Classification: LCC GB5014 .A24 2023 | DDC 904/.5—dc23 LC record available at https://lccn.loc.gov/2021026407

The Internet addresses listed in the text were accurate at the time of publication. The inclusion of a website does not indicate an endorsement by the authors or McGraw Hill LLC, and McGraw Hill LLC does not guarantee the accuracy of the information presented at these sites.

mheducation.com/highered











About the Author



Patrick L. Abbott Patrick Abbott is a native San Diegan. Pat earned his MA and PhD degrees in geology at the University of Texas at Austin. He benefited greatly from the depth and breadth of the faculty in the Department of Geological Sciences at Austin; this was extended by their requirement to take five additional graduate courses outside the department. Developing interests in many topics helped lead to writing this textbook.

Pat's research has concentrated on the Mesozoic and Cenozoic sedimentary rocks of the southwestern United States and northwestern Mexico. Studies have focused on reading the history stored within the rocks—depositional environments, provenance, paleoclimate, palinspastic reconstructions, and high-energy processes.

Pat has long been involved in presenting Earth knowledge to the public, primarily through TV news. He has produced award-winning videos for TV broadcast. He was one of the main cast members in the TV series *The Real Gilligan's Island* on TBS, *Serial Killer Earth* on H2 (The History Channel 2), and *So You Think You'd Survive* on The Weather Channel. During part of each year, Pat works as a Smithsonian lecturer visiting all continents and oceans.









Brief Contents

Preface xv

Prologue: Energy Flows 1

- 1 Natural Disasters and the Human Population 5
- 2 Internal Energy and Plate Tectonics 23
- 3 Earthquake Geology and Seismology 47
- 4 Plate Tectonics and Earthquakes 77
- 5 Earthquakes Throughout the United States and Canada 108
- 6 Volcanic Eruptions: Plate Tectonics and Magmas 144
- 7 Volcano Case Histories: Killer Events 176
- 8 Tsunami Versus Wind-Caused Waves 205
- 9 External Energy Fuels Weather and Climate 231
- 10 Tornadoes, Lightning, Heat, and Cold 255
- 11 Hurricanes 288
- 12 Climate Change 320
- **13** Floods 358
- **14** Wildfire 389
- **15** Mass Movements 417
- 16 Coastal Processes and Hazards 453
- **17** Impacts with Space Objects 467
- **18** The Great Dyings 493

Geologic Time Table and Table of Conversions A-1

Glossary G-1

Index I-1







Contents

Preface xv

Prologue: Energy Flows 1

Processes of Construction versus Destruction 3

Terms to Remember 4

CHAPTER 1

Natural Disasters and the Human Population 5

Great Natural Disasters 6

Human Fatalities and Economic Losses

in Natural Disasters 6

The Role of Government in Natural-Disaster

Death Totals 6

Human Responses to Disaster 8

Economic Losses from Natural Disasters 8

Natural Hazards 8

Popocatépetl Volcano, Mexico 9

Magnitude, Frequency, and Return Period 9

Role of Population Growth 10

Overview of Human Population 10

The Power of an Exponent on Growth 11

The Past 10,000 Years of Human History 12

The Human Population Today 12

Future World Population 13

Demographic Transition 14

Side Note: Interest Paid on Money: An Example

of Exponential Growth 15

Urbanization and Earthquake Fatalities 16

A Classic Disaster: Influenza (Flu) Pandemic of 1918 17

Disease Pandemics 17

Carrying Capacity 19

Summary 21; Terms to Remember 21; Questions for Review 22; Questions for Further Thought 22

CHAPTER 2

Internal Energy and Plate Tectonics 23

Origin of the Sun and Planets 24

Impact Origin of the Moon 25

Earth History 25

The Layered Earth 26

Side Note: Mother Earth 27

Behavior of Materials 27

Side Note: Volcanoes and the Origin of the Ocean,

Atmosphere, and Life 28

Isostasy 28

Internal Sources of Energy 30

Impact Energy and Gravitational Energy 30

Radioactive Isotopes 30

In Greater Depth: Radioactive Isotopes 32

Age of Earth 32

In Greater Depth: Radioactivity Disasters 33

Plate Tectonics 33

Development of the Plate Tectonics Concept 34

In Greater Depth: Earth's Magnetic Field 36

Magnetization of Volcanic Rocks 37

The Grand Unifying Theory 43

How We Understand Earth 45

Uniformitarianism 45

Summary 45; Terms to Remember 46; Questions for Review 46; Questions for Further Thought 46

CHAPTER 3

Earthquake Geology and Seismology 47

A Classic Disaster: The Lisbon Earthquake of 1755 48

Understanding Earthquakes 48

Faults and Geologic Mapping 49

Types of Faults 51



Dip-Slip Faults 51 Strike-Slip Faults 52

Transform Faults 54

Development of Seismology 56

Waves 56

Seismic Waves 57

Body Waves 57

Seismic Waves and Earth's Interior 57

Surface Waves 57

Sound Waves and Seismic Waves 58

In Greater Depth: Seismic Waves from Nuclear Bomb

Blasts versus Earthquakes 60

Locating the Source of an Earthquake 60

Magnitude of Earthquakes 61

Richter Scale 61

Other Measures of Earthquake Size 63

Foreshocks, Mainshock, and Aftershocks 63

In Greater Depth: The Search for Tiny Earthquakes 64

Magnitude, Fault-Rupture Length, and Seismic-

Wave Frequencies 64

In Greater Depth: F = ma 65

Ground Motion During Earthquakes 65

Acceleration 65

Periods of Buildings and Responses of

Foundations 65

In Greater Depth: What to Do Before and During

an Earthquake 66

Earthquake Intensity—What We Feel During

an Earthquake 66

Mercalli Scale Variables 67

A Case History of Mercalli Variables:

The San Fernando Valley, California,

Earthquake of 1971 68

Learning from the Past 70

Building in Earthquake Country 71

Shear Walls and Bracing 71

Retrofit Buildings, Bridges, and House

Construction 72

Base Isolation 73

Summary 74; Terms to Remember 75; Questions for Review 75; Questions for Further Thought 76;

Disaster Simulation Game 76

CHAPTER 4

Plate Tectonics and Earthquakes 77

Tectonic-Plate Edges and Earthquakes 78

Spreading-Center Earthquakes 79

Iceland 79

Red Sea and Gulf of Aden 80

Gulf of California 82

Convergent Zones and Earthquakes 83

Subduction-Zone Earthquakes 83

Japan, 2011: Stuck Segments of Subducting

Plate 84

A Classic Disaster: The Tokyo Earthquake of 1923 85

Indonesia, 2004: One Earthquake Triggers

Others 85

Mexico City, 1985: Long-Distance Destruction 85

Chile, 1960: The Biggest One 88

Alaska, 1964: Second Biggest One 88

Pacific Northwest: The Upcoming Earthquake 89

Continent-Continent Collision Earthquakes 90

Nepal, China, Pakistan, and India, 2015, 2008,

2005, and 2001: Continent Collision Kills 91

China, 1556: The Deadliest Earthquake 92

The Arabian Plate 92

Continent-Continent Collision Earthquakes 92

Transform-Fault Earthquakes 92

Transform-Fault Earthquakes 94

Haiti, 2010: Earthquakes Don't Kill, Buildings

Do 94

Turkey, 1999: Serial Earthquakes 95

San Andreas Fault Tectonics and Earthquakes 96

A Classic Disaster: The San Francisco

Earthquake of 1906 102

Bay Area Earthquakes—Past and Future 102

Summary 106; Terms to Remember 106;

Questions for Review 106; Questions for Further

Thought 107

CHAPTER 5

Earthquakes Throughout the United States and Canada 108

How Faults Work 109

Elastic Rebound 109

Newer View 109

Thrust-Fault Earthquakes 112

Virginia, 2011: Ancient Faults Can

Reactivate 112

Northridge, California, 1994: Compression

at the Big Bend 112

Seattle, Washington 113







Normal-Fault Earthquakes 114

Puget Sound, Washington, 1949, 1965, 2001: Subducting Plates Can Crack 114

Neotectonics and Paleoseismology 114

Earthquake Prediction 116

Long-Term Forecasts 116

Side Note: Perils of Prediction: Scientists

on Trial 118

Short-Term Forecasts 118

Things to Keep in Mind 119

Early Warning Systems 119

Shakemaps 120

Did You Feel It? 120

Great Shakeout Events 120

Earthquake Losses for the United States 120

Human-Triggered Earthquakes 120

Pumping Fluids Underground 122

Dam Earthquakes 123

Bomb Blasts 123

Earthquakes in the United States and

Canada 123

Western North America: Plate Boundary–Zone

Earthquakes 126

Western Great Basin: Eastern California, Western

Nevada 126

Intermountain Seismic Belt: Utah, Idaho,

Wyoming, Montana 129

Rio Grande Rift: New Mexico, Colorado,

Westernmost Texas, Mexico 131

Intraplate Earthquakes: "Stable" Central United States 132

New Madrid, Missouri, 1811-1812 132

Reelfoot Rift: Missouri, Arkansas, Tennessee,

Kentucky, Illinois 134

Ancient Rifts in the Central United States 135

Intraplate Earthquakes: Eastern North

America 137

New England 137

St. Lawrence River Valley 138

Charleston, South Carolina, 1886 138

Earthquakes and Volcanism in Hawaii 140

Earthquake in 1975 141

Earthquakes in 2006 141

Earthquake in 2018 141

Summary 141; Terms to Remember 142; Questions for Review 142; Questions for Further Thought 142

CHAPTER 6

Volcanic Eruptions: Plate Tectonics and Magmas 144

How We Understand Volcanic Eruptions 145 Plate-Tectonic Setting of Volcanoes 145

A Classic Disaster: Eruption of Mount

Vesuvius, 79 CE 147

Chemical Composition of Magmas 148

Viscosity, Temperature, and Water Content

of Magmas 149

In Greater Depth: Minerals and Volcanic Rocks 150

Plate-Tectonic Setting of Volcanoes Revisited 152

How a Volcano Erupts 153

Anatomy of a Magma Body 154

Eruption Styles and the Role of Water

Content 154

Some Volcanic Materials 155

The Three Vs of Volcanology: Viscosity,

Volatiles, Volume 157

Shield Volcanoes: Low Viscosity, Low Volatiles,

Large Volume 157

Side Note: How a Geyser Erupts 158

Flood Basalts: Low Viscosity, Low Volatiles, Very

Large Volume 160

In Greater Depth: Volcanic Explosivity Index (VEI) 161

Scoria Cones: Medium Viscosity, Medium

Volatiles, Small Volume 162

Stratovolcanoes: High Viscosity, High Volatiles,

Large Volume 162

Lava Domes: High Viscosity, Low Volatiles,

Small Volume 165

Side Note: British Airways Flight 9 166

Calderas: High Viscosity, High Volatiles, Very

Large Volume 166

In Greater Depth: Hot Spots 172

A Classic Disaster: Santorini and the Lost Island

of Atlantis 173

Summary 174; Terms to Remember 174;

Questions for Review 174; Questions for Further

Thought 175

CHAPTER 7

Volcano Case Histories: Killer Events 176

Volcanism at Spreading Centers 177 Iceland 177







Volcanism at Subduction Zones 178

Cascade Range, Pacific Coast of United States and Canada 178

In Greater Depth: Rapid Assembly and Rise of Magma 188

Volcanic Processes and Killer Events 189
The Historic Record of Volcano Fatalities 189
Pyroclastic Eruptions 189

A Classic Disaster: Mont Pelée, Martinique, 1902 191 Tsunami 193

Lahars 193

Side Note: Death at Ashfall, Nebraska 196

Debris Avalanches 196 Indirect—Famine 197

Gas 197

Lava Flows 199

VEIs of Some Killer Eruptions 199

Volcano Monitoring and Warning 200

Long Valley, California 200

Mount Pinatubo, Philippines, 1991 202

Signs of Impending Eruption 202

Volcano Observatories 203

Summary 204; Terms to Remember 204; Questions for Review 204; Questions for Further Thought 204

CHAPTER 8

Tsunami Versus Wind-Caused Waves 205

Japanese Tsunami, 11 March 2011 206
Tsunami Travel Through the Pacific Ocean 206
Land Subsidence 207

Side Note: Fukushima Daiichi Nuclear Disaster 208 British Columbia, Washington, Oregon, 26 January 1700 208

Wind-Caused Waves 208 Why a Wind-Blown Wave Breaks 209

Rogue Waves 209

In Greater Depth: Deep-Water Wave Velocity, Length,

Period, and Energy 211

Tsunami 211

Tsunami versus Wind-Caused Waves 212

A Classic Disaster: The Chile Tsunami

of 1868 215

Tsunami at the Shoreline 215

Earthquake-Caused Tsunami 216

Indian Ocean 26 December 2004 217

Alaska, 1 April 1946: First Wave Biggest 218

Chile, 22 May 1960: Third Wave Biggest 219

Alaska, 27 March 1964: Fifth Wave Biggest 220

Volcano-Caused Tsunami 220

Krakatau, Indonesia, 26-27 August 1883 220

Landslide-Caused Tsunami 221

Volcano Collapses 221

Earthquake-Triggered Mass Movements 222

In Bays and Lakes 224

Seiches 225

Hebgen Lake, Montana, 17 August 1959 225

Tsunami and You 226

Simeulue Island, Indonesia, 26 December 2004 226

Nicaragua, 1 September 1992 226

Humans Can Increase the Hazard 227

Tsunami Warnings 227

Summary 229; Terms to Remember 229; Questions for Review 229; Questions for Further Thought 229; Disaster Simulation Game 230

CHAPTER 9

External Energy Fuels Weather and Climate 231

External Sources of Energy 232

The Sun 232

Solar Radiation Received by Earth 233

Outgoing Terrestrial Radiation 234

Greenhouse Effect 234

Albedo 234

Convection and Conduction 234

The Hydrologic Cycle 236

In Greater Depth: Water—The Most Peculiar

Substance on Earth? 237

Water and Heat 237

Water Vapor and Humidity 237

Latent Heat 238

Adiabatic Processes 238

Lapse Rates 239

Differential Heating of Land and Water 239

Energy Transfer in the Atmosphere 239

Energy Transfer in the World Ocean 241

Layering of the Lower Atmosphere 241

Temperature 241

Pressure 242

Winds 242

Pressure Gradient Force 242

viii Contents



In Greater Depth: Coriolis Effect 243

Rotating Air Bodies 243

General Circulation of the Atmosphere 244

Low Latitudes 245

Middle Latitudes 246

High Latitudes 249

Observed Circulation of the Atmosphere 250

General Circulation of the Oceans 250

Surface Circulation 250

Deep-Ocean Circulation 251

Summary 253; Terms to Remember 253; Questions for Review 253; Questions for Further Thought 254

CHAPTER 10

Tornadoes, Lightning, Heat, and Cold 255

Severe Weather 256

Winter Storms 257

Cold 257

Precipitation 257

Bomb Cyclone 257

Nor'easters 258

In Greater Depth: Doppler Radar 259

Blizzards 260

Ice Storms 260

Lake-Effect Snow 260

How Thunderstorms Work 261

Lifting of Air 262

Air-Mass Thunderstorms 262

Severe Thunderstorms 263

Supercells 264

Thunderstorms in North America 264

In Greater Depth: Downbursts: An Airplane's

Enemy 265

Heavy Rains and Flash Floods 266

Hail 267

Derechos 268

Tornadoes 269

Tornadoes in 2011 269

How Tornadoes Form 269

Regional Scale 270

Supercell Thunderstorm Scale 270

Vortex Scale 272

Tornadoes in the United States and Canada 273

A Classic Disaster: The Tri-State Tornado of 1925 276

Tornado Outbreaks 276

Tornadoes and Cities 278

How a Tornado Destroys a House 279

Tornado Safety 279

Safe Rooms 279

Lightning 279

How Lightning Works 281

Don't Get Struck 281

Heat 284

Heat Wave in Chicago, July 1995 284

City Weather 285

European Heat Waves, 2003 and 2010 285

Summary 286; Terms to Remember 287; Questions for Review 287; Questions for Further Thought 287

CHAPTER 11

Hurricanes 288

Hurricanes 289

How a Hurricane Forms 290

How a Hurricane Works 290

Eyewall and Eye 292

Tornadoes within Hurricanes 292

Eyewall Replacement Cycle 292

Energy Flow in a Hurricane 293

Hurricane Energy Release 293

Hurricane Transition to Post-Tropical

Cyclone 293

Hurricane Origins 294

North Atlantic Ocean Hurricanes 296

Cape Verde-Type Hurricanes 296

Caribbean Sea and Gulf of Mexico-Type

Hurricanes 299

Hurricane Forecasts 300

Rapid Intensification 302

How Hurricanes Get Their Names 303

Accumulated Cyclone Energy 303

Hurricane Damages 304

The Big Three Hazards 305

Storm-Surge Hazards 306

Inland Flooding 307

A Classic Disaster: The Galveston Hurricane

of 1900 309

Wind 310

Preparation for Hurricanes 310

Hurricane Katrina, August 2005 310

The Evacuation Dilemma 312







Reduction of Hurricane Damages 313 Land-Use Planning 314

In Greater Depth: How to Build a Home Near

the Coastline 315

Global Rise in Sea Level 315

Hurricanes and the Pacific Coastline 315

Hurricane Iniki, September 1992 316

Cyclones and Bangladesh 316

Summary 318; Terms to Remember 318; Questions for Review 318; Questions for Further Thought 319; Disaster Simulation Game 319

CHAPTER 12

Climate Change 320

Early Earth Climate—An Intense Greenhouse 321 Climate History of Earth: Timescale in Millions of Years 322

In Greater Depth: Equilibrium Between Tectonics, Rock Weathering, and Climate 323

Late Paleozoic Ice Age 324

Late Paleocene Torrid Age 324

In Greater Depth: Oxygen Isotopes and Temperature 326

Late Cenozoic Ice Age 326

Glacial Advance and Retreat: Timescale

in Thousands of Years 326

The Last Glacial Maximum 328

Climate Variations: Timescale in Hundreds

of Years 329

Shorter-Term Climate Changes: Timescale

in Multiple Years 331

El Niño 331

La Niña 333

Marine Heat Waves 334

Volcanism and Climate 334

Volcanic Climate Effects 336

In Greater Depth: The Mayan Civilization and Climate

Change 337

The Past Thousand Years 337

The 20th Century 339

Solar Energy Variation 339

Side Note: Stradivari Violins 340

Radiative Forcing 340

In Greater Depth: When Did Humans Begin Adding

to Greenhouse Warming? 341

Greenhouse Gases and Aerosols 341

Water Vapor 341

Carbon Dioxide (CO₂) 342

Methane (CH₄) 342

Nitrous Oxide (N₂O) 343

Ozone (O₃) 343

Chlorofluorocarbons (CFCs) 343

20th-Century Greenhouse Gas Increases 343

Aerosols 343

The 21st Century 343

Global Climate Models 344

Drought and Famine 346

Ice Melting 347

In Greater Depth: Tipping Points 349

Sea-Level Rise 350

In Greater Depth: Lag Times 351

Ocean Changes 352

Signs of Change 353

Mitigation Options 353

Controlling Co₂ Content of Atmosphere 354

Managing Incoming Solar Radiation 354

Fast-Action Strategies 355

Climate-Driven Human Migrations 355

Summary 355; Terms to Remember 356; Questions for Review 357; Questions for Further Thought 357

CHAPTER 13

Floods 358

How Rivers and Streams Work 359

Side Note: A Different Kind of Killer Flood 360

The Equilibrium Stream 360

In Greater Depth: Stream Velocity Profile 362

Graded-Stream Theory 363

The Floodplain 363

Side Note: Feedback Mechanisms 364

Flood Frequency 364

Florence, Italy, 1333 and 1966 364

Flood-Frequency Curves 364

In Greater Depth: Constructing Flood-Frequency

Curves 366

Flood Styles 367

Flash Floods vs. Regional Floods 368

Flash Floods 368

Regional Floods 372

Societal Responses to Flood Hazards 378

Dams, Reservoirs, and Natural Storage Areas 378

Levees 379





Sandbagging 380 Forecasting 380 Zoning and Land Use 380 Insurance 381 Presidential Disaster Declarations 381 Urbanization and Floods 381 Hydrographs 381 Flood Frequencies 381 Channelization 382 The Biggest Floods 384 Ice-Dam Failure Floods 385 Summary 387; Terms to Remember 388; Questions for Review 388; Questions for Further Thought 388; Disaster Simulation Game 388 **CHAPTER 14** Wildfire 389 Fire 390 In Greater Depth: Fire and Life 391 What Is Fire? 391 The Fire Triangle 391 A Classic Disaster: The Burning of Rome, 64 CE 392 The Fuels of Fire 392 Grasses 393 Shrubs 393 Forests 394 Houses 394 The Causes of Wildfires 395 Lightning Siege, California 2020 395 The Stages of Fire 395 The Spread of Fire 396 **Fuel 397** Topography 397 Fire Behavior 397 Wind 398 Fire Weather and Winds 398

Local Winds 399

Region, 1871 400

Downslope Winds 399

The End of Normal 404

Wind and Fire in California 401

The Western United States 404

Smoke Effects on Human Health 404

A Classic Disaster: Wind and Fire in the Great Lakes

Home Design and Fire 405 Side Note: The Winds of Madness 406 How Well Have We Learned? 407 Fire Suppression 407 Yellowstone National Park Wildfire 408 California versus Baja California: Pay Now or Pay Later 409 The Western and Southern United States in 2000 411 Prescribed Fires 412 Wildfires in Australia 413 The Similarities of Fire and Flood 414 Summary 414; Terms to Remember 415; Questions for Review 415; Questions for Further Thought 415; Disaster Simulation Game 416 **CHAPTER 15** Mass Movements 417 The Role of Gravity in Mass Movements 418 Creep 419 External Causes of Slope Failures 420 In Greater Depth: Energy, Force, Work, Power, and Heat 421 Water in Its External Roles 422 Internal Causes of Slope Failures 422 Inherently Weak Materials 422 Water in Its Internal Roles 423 In Greater Depth: Analysis of Slope Stability 425 Decreases in Cohesion 425 A Classic Disaster: Vaiont Landslide, Italy, 1963 426 Adverse Geologic Structures 427 Triggers of Mass Movements 428 Classification of Mass Movements 428 Falls 428 Yosemite National Park, California 429

Slides 429

Rotational Slides 430

Translational Slides 432

Flows 434

Portuguese Bend, California, Earthflow 434 La Conchita, California, Slump and Debris Flows, 1995 and 2005 436

Long-Runout Debris Flows 436

Snow Avalanches 441

Submarine Mass Movements 443

Contents



Mitigation 444

Reshaping Topography 444

Strengthening Slopes 444

Draining Water 445

Controlling Erosion 446

Subsidence 447

Catastrophic Subsidence 447

Slow Subsidence 448

In Greater Depth: How to Create a Cave 449

Summary 451; Terms to Remember 452; Questions for Review 452; Questions for Further

Thought 452

CHAPTER 16

Coastal Processes and Hazards 453

Sand 454

Summer versus Winter Beaches 455

Waves 455

Rip Currents 455

Side Note: Shark-Attack Deaths 457

Wave Refraction 457

Longshore Drift 457

In Greater Depth: Gravity and Tides 459

Tides 459

Tidal Bores 460

King Tides 460

Coastal-Control Structures 460

Seawalls 460

Cliff Armoring 461

Groins and Jetties 462

Breakwaters 462

Massive Structures in Future 463

Mother Nature at Work 463

Mangrove Forests 463

Side Note: You Can Never Do Just One Thing 464

Coral-Algal Reefs 464

Summary 465; Terms to Remember 466; Questions for Review 466; Questions for Further Thought 466

CHAPTER 17

Impacts with Space Objects 467

Energy and Impacts 468

Impact Scars 469

Sources of Extraterrestrial Debris 470

Asteroids 470

xii Contents

Side Note: Dwarf Planets 471

Comets 473

In Greater Depth: Insights from Spacecraft Landing

on a Comet 474

Rates of Meteoroid Influx 475

Cosmic Dust 475

In Greater Depth: Shoemaker-Levy 9 Comet Impacts

on Jupiter 476

Shooting Stars 477

Meteorites 477

The Crater-Forming Process 478

Crater-Forming Impacts 479

Meteor Crater, Arizona 480

Impact Origin of Chesapeake Bay 482

The End Cretaceous Impact 483

Evidence of the End Cretaceous Impact 484

Site of the End Cretaceous Impact 484

Size and Velocity of Impactor 484

Angle of Impact 484

Problems for Life from the End Cretaceous

Impact 484

Biggest Impact Events of the 20th and 21st

Centuries 485

Tunguska, Siberia, 1908 485

Big "Near Events" 487

Frequency of Large Impacts 487

In Greater Depth: Coronal Mass Ejections 488

Lifetime Risks of Impact 488

Prevention of Impacts 490

Summary 491; Terms to Remember 491; Questions for Review 491; Questions for Further Thought 492

CHAPTER 18

The Great Dyings 493

Fossils 494

Early Understanding of Extinctions

and Geologic Time 496

Brief History of Life 498

Species and the Fossil Record 499

The Tropical Reef Example 500

Mass Extinctions During Phanerozoic Time 501

Possible Causes of Mass Extinctions 502

Plate-Tectonic Causes 503

Volcanic Causes 504

Climate Change Causes 505

Ocean Composition Causes 505



Extraterrestrial Causes 505 Biologic Causes 505

Multiple Causes of Mass Extinction 507

Examples of Mass Extinctions 507

Close of Permian Period (Ended 252 Million Years Ago) 507

In Greater Depth: Life Moves onto Land During Paleozoic Time 508

Close of Cretaceous Period (Ended 66 Million Years Ago) 509

In Greater Depth: Dinosaurs Dominate the Land

During Mesozoic Time 510 Living Fossils 512

In Greater Depth: The Rise of Mammals and Birds

During Cenozoic Time 513

The Current Mass Extinction 514

In Greater Depth: La Brea Tar Pits, Metropolitan Los

Angeles 516

Australia 516

Side Note: The Rewilding of North America 517

Madagascar and New Zealand 517

In Greater Depth: De-extinction: The Resurrection

of Extinct Species 518

Summary 518; Terms to Remember 519; Questions for Review 519; Questions for Further

Thought 519

Geologic Time Table and Table of Conversions A-1

Glossary G-1 Index I-1











Preface

Why Study Natural Disasters?

Natural disasters occur every day and affect the lives of millions of people each year. Many students have been affected by earthquakes or tornadoes or hurricanes or floods or landslides or wildfires or other events. They are interested in lectures and images that explain these processes, and lively discussions commonly ensue.

During decades of teaching courses at San Diego State University, I found that students have an innate curiosity about "death and destruction"; they want to know why natural disasters occur. Initiation of a Natural Disasters course led to skyrocketing enrollments that exceeded 5,000 students per year. Some of these experiences are described in a *Journal of Geoscience Education* article by Pat Abbott and Ernie Zebrowksi [v 46 (1998), pp. 471–75].

Themes and Approach

This textbook focuses on explaining how the normal processes of the Earth concentrate their energies and deal heavy blows to humans and their structures. The following themes are interwoven throughout the book:

- Energy sources underlying disasters
- Plate tectonics
- · Climate change
- Earth processes operating in rock, water, and atmosphere
- Significance of geologic time
- Complexities of multiple variables operating simultaneously
- Detailed and interesting case histories
- · Population growth

New to This Edition

- Chapter 1: Updated concepts; revised In Greater Depth: 1918 Flu; added section on zoonotic diseases; added discussion of coronavirus pandemic, 2020-2021; updated tables and line art.
- Chapter 2: Incorporated new understanding of origins of Earth and Moon; revised discussion of isostasy.

- Chapter 3: Incorporated new understanding of seismic waves and fault behavior; expanded discussion of stress vs. strain; added discussion of power laws; discussed significance of very small earthquakes; revised In Greater Depth: F = ma.
- Chapter 4: Updated tables; revised map figures.
- Chapter 5: Added discussion of Ridgecrest earthquakes of 2019 to show new pattern of fault movements, including new figure; added FEMA annualization of earthquake losses in U.S. with new figure; revised discussion of Oklahoma earthquake magnitude and frequency; updated discussion of earthquake warning systems; added Hawaii 2018 magnitude 6.9 earthquake.
- Chapter 6: Added new section on anatomy of a magma body, with new line art; expanded discussion of Naples/ Campi Flegrei eruption potential; added table of Vesuvius eruption history; expanded discussion of VEI; added discussion of supervolcanoes; added photo of Mount Redoubt.
- Chapter 7: Discussed likelihood of Mt. St. Helens future eruptions; added White Island, New Zealand eruption, which killed 24 tourists in 2019; updated discussion of Long Valley, California, eruption potential.
- Chapter 8: Updated tables; expanded discussion of breaking waves; added Krakatau Volcano collapse in 2018.
- Chapter 9: Expanded section on atmospheric rivers, including new figure and table; expanded discussion of polar vortex, including new figures.
- Chapter 10: Updated tables and figures; expanded discussion of bomb cyclone; compared body heat and wetbulb temperatures; replaced derechos case history with Branson, Missouri, 2018 tragedy; mentioned huge Iowa derecho 2020; new table on lightning deaths.
- Chapter 11: Updated tables; new section on wind, with table; rewrote section on hurricane forecasting; emphasized The Big 3: storm surge, wind, inland flooding; new section The Slowdown, with new figure.
- Chapter 12: Updated tables; expanded discussion of CO₂; new section on marine heat waves; expanded discussion of volcanism and climate change; new section on climate sensitivity; added discussion of methane



bursts in Siberia; added discussion of Arctic wildfires; new section Climate-Driven Human Migrations.

- Chapter 13: Updated tables; expanded discussion of atmospheric rivers, with new image.
- Chapter 14: New discussion of western North America megafires, 2020; new section on lightning siege wildfires; new section on firenadoes; expanded discussion of effects of smoke on health; new section, The End of Normal; properties at risk in U.S. states, with tables; expanded discussion of prescribed fires; updated Australia bushfires.
- Chapter 15: New table on U.S. snow avalanche fatalities by state; expanded discussion of subsidence to include Corvette Museum in Kentucky; new paragraph on insurance choices.
- Chapter 16: Expanded discussion of the tragedy of the commons; new section on king tides; updated discussion of rip currents; new section on mangrove forests.
- Chapter 17: Updated statistics; new images.
- Chapter 18: Extensive revisions and updates of chapter; new or revised sections on the concept of species, keystone species, duration of Permian extinction, post-Cretaceous life recovery, de-extinction; new In Greater Depth boxes follow rise and fall of amphibians, dinosaurs, mammals, and birds.

Acknowledgments

I am deeply appreciative of the help given by others to make this book a reality. The photograph collection in the book is immeasurably improved by the aerial photographs generously given by the late John S. Shelton, the greatest geologist photographer of them all. Please see John's classic book *Geology Illustrated*.

The quality of the book was significantly improved by the insights provided by comments from the following reviewers:

Austin Community College, Kusali Gamage
Baylor University, John Dunbar
Bellevue College, Ian Walker
Bloomsburg University of Pennsylvania, Benjamin Franek
California State Polytechnic University-Pomona,
Jon Nourse

California State University, Chico, Susan G. Riggins California State University, Los Angeles, Richard W. Hurst California State University, Sacramento, Lisa Hammersley California State University—Fullerton, Stephen I. Wareham California State University—Los Angeles,

Hassan Rezaie Boroon

Chandler–Gilbert Community College, John Dassinger

Colby College, James Rodger Fleming

College of Southern Idaho, Shawn P. Willsey

Columbus State Community College, Jeffery G. Richardson

Dartmouth College, Leslie Sonder
Delaware County Community College, William J. Dyke, Jr.
Erie Community College, Buffalo State College,
Karen S. Wehn

Fairleigh Dickinson University, Edward Catanzaro Fort Lewis College, David Gonzales Grand Valley State University, Patrick Colgan Illinois State University, Robert S. Nelson Indiana University, Bingming Shen-Tu Lindenwood University, Sandra Allen Manchester Community College, Eszter Samodai Minnesota State University-Mankato, Cecil S. Keen Morehead State University, Eric Jerde Northeastern University, Jennifer Cole Northeastern University, Langdon D. Clough North Hennepin Community College, John Dooley Penn State University-Altoona, Timothy J. Dolney Pennsylvania State University, Kevin P. Furlong Radford University, Jonathan Tso Rio Hondo Junior College, Michael Forrest Salisbury University, Brent R. Skeeter San Diego State University, Jim Rickard San Diego State University, Victor E. Camp San Francisco State University, Bridget James San Francisco State University, Mary Leech San Francisco State University, Oswaldo Garcia Sonoma State University, Terry Wright Southeast Arkansas University, Steven Sumner Southeast Missouri State University, Ernest L. Kern St. Cloud State University, Alan Srock St. Cloud State University, Jean L. Hoff SUNY-Stony Brook, Christiane Stidham Temple University, Jesse Thornburg Temple University, Tim Davis Texas State University, Philip Suckling The Arizona Geological Survey, Michael Conway The Ohio State University, Michael Barton Tulane University, Stephen A. Nelson University at Albany, Michael G. Landin University of British Columbia, Roland Stull University of California, Santa Barbara, Robin Matoza University of California, Santa Cruz, Thorne Lay University of California-Davis, John F. Dewey University of California-Riverside, Peter Sadler University of California-San Diego, Gabi Laske University of California-Santa Barbara, Cathy Busby University of Colorado, Alan Lester University of Colorado-Boulder, Charles R. Stern University of Colorado-Colorado Springs, Paul K. Grogger University of Illinois at Urbana-Champaign,

Wang-Ping Chen
University of Iowa, David W. Peate
University of Kansas, David Braaten
University of Kansas, Don Steeples
University of Kentucky–Lexington, Kevin Henke

vi Preface



University of Michigan, Youxue Zhang University of Nebraska, Nathan Eidem University of Nebraska at Kearney, Jeremy S. Dillon University of Nebraska-Kearney, A. Steele Becker University of Nebraska-Kearney, Jean Eichhorst University of Nebraska-Kearney, Stanley Dart University of Nebraska-Kearney, Vijendra K. Boken University of North Carolina at Chapel Hill, Melissa Hudley University of North Carolina-Greensboro, John Hidore University of Oklahoma, Barry Weaver University of Oklahoma, Judson Ahern University of Portland, Robert Butler University of Southern California, John P. Wilson University of Wisconsin-LaCrosse, George Hupper Utah State University, Susan K. Morgan Washington University-St. Louis, Carol Prombo Yale University, David Bercovici

Special thanks to the following individuals who wrote and/or reviewed learning goal-oriented content for **LearnSmart**.

California State University–Sacramento, Lisa Hammersley Northern Arizona University, Sylvester Allred Roane State Community College, Arthur C. Lee

I sincerely appreciate the talents and accomplishments of the McGraw Hill professionals in Dubuque who took my manuscript and produced it into this book. For the shortcomings that remain in the book, I alone am responsible. I welcome all comments, pro and con, as well as suggested revisions.

Pat Abbott professor_pat_abbott@yahoo.com

Preface xvii





Instructors: Student Success Starts with You

Tools to enhance your unique voice

Want to build your own course? No problem. Prefer to use an OLC-aligned, prebuilt course? Easy. Want to make changes throughout the semester? Sure. And you'll save time with Connect's auto-grading too.

65% **Less Time** Grading



Laptop: McGraw Hill; Woman/dog: George Doyle/Getty Image

Study made personal

Incorporate adaptive study resources like SmartBook® 2.0 into your course and help your students be better prepared in less time. Learn more about the powerful personalized learning experience available in SmartBook 2.0 at

www.mheducation.com/highered/connect/smartbook

Affordable solutions, added value



Make technology work for you with LMS integration for single sign-on access, mobile access to the digital textbook, and reports to quickly show you how each of your students is doing. And with our Inclusive Access program you can provide all these tools at a discount to your students. Ask your McGraw Hill representative for more information.

Solutions for your challenges



A product isn't a solution. Real solutions are affordable, reliable, and come with training and ongoing support when you need it and how you want it. Visit www .supportateverystep.com for videos and resources both you and your students can use throughout the semester.





Students: Get Learning that Fits You

Effective tools for efficient studying

Connect is designed to help you be more productive with simple, flexible, intuitive tools that maximize your study time and meet your individual learning needs. Get learning that works for you with Connect.

Study anytime, anywhere

Download the free ReadAnywhere app and access your online eBook, SmartBook 2.0, or Adaptive Learning Assignments when it's convenient, even if you're offline. And since the app automatically syncs with your Connect account, all of your work is available every time you open it. Find out more at www.mheducation.com/readanywhere

"I really liked this app—it made it easy to study when you don't have your textbook in front of you."

- Jordan Cunningham, Eastern Washington University



Everything you need in one place

Your Connect course has everything you need—whether reading on your digital eBook or completing assignments for class, Connect makes it easy to get your work done.

Calendar: owattaphotos/Getty Images

Learning for everyone

McGraw Hill works directly with Accessibility Services Departments and faculty to meet the learning needs of all students. Please contact your Accessibility Services Office and ask them to email accessibility@mheducation.com, or visit www.mheducation.com/about/accessibility

for more information.

Top: Jenner Images/Getty Images, Left: Hero Images/Getty Images, Right: Hero Images/Getty Images







CHAPTER 1



Natural Disasters and the Human Population

Mankind was destined to live on the edge of perpetual disaster. We are mankind because we survive.

—James A. Michener, 1978, Chesapeake, Random House



The world population of humans continues to increase exponentially. Photo of shopping area in New Delhi, India.

donyanedomam/123RF

LEARNING OUTCOMES

The human population is growing rapidly. Natural disasters are causing great numbers of deaths and economic losses. After studying this chapter you should

- recognize the differences between a natural hazard, a natural disaster, and a great natural disaster.
- be familiar with the processes that cause the deadliest natural disasters.
- understand the relationship between frequency and magnitude of natural disasters.
- know the size of the human population.
- understand the significance of exponential growth.
- recognize the demographic transition of human populations.
- understand the deadliness of pandemics.
- be able to explain the concept of carrying capacity.

OUTLINE

- Great Natural Disasters
- Human Fatalities and Economic Losses in Natural Disasters
- Natural Hazards
- Overview of Human Population
- Future World Population
- Carrying Capacity





n 2018 there were 315 **natural disasters**; they killed a total of 11,804 people. The worst killer events were **earthquakes**, **tsunami**, volcanoes, floods, heat waves, and wildfires. The 11 deadliest events are listed in table 1.1; they occurred in seven different months and nine different countries, mostly in Asia. As horrible as the 2018 death total is, it is markedly less than the average of 67,572 killed each year in the preceding decade (2008 to 2017). All these disasters were the result of natural processes operating at high **energy** levels for brief times in restricted areas.

Great Natural Disasters

The Japan earthquake and tsunami in 2011, the Haiti earthquake in 2010, and the Myanmar cyclone and China earthquake in 2008 combined to kill almost 500,000 people. They are examples of great natural disasters: these events so overwhelm regions that international assistance is needed to rescue and care for people, clean up the destruction, and begin the process of reconstruction. Great natural disasters commonly kill thousands of people, leave hundreds of thousands homeless, and overwhelm the regional economy.

Today, in earthquake-active areas of the world, several hundred million people live in buildings that will collapse during a strong earthquake. An earthquake killing more than 100,000 people could happen any day in Teheran, Iran; in Istanbul, Turkey; or in other large cities. Today, people by the millions are moving to the ocean shores, where they can be hit by tsunami, hurricanes, and floods. We need to learn how to build disaster-resistant communities to lessen the human fatalities and economic losses resulting from natural disasters.

Human Fatalities and Economic Losses in Natural Disasters

The 40 deadliest disasters in the 47-year period from 1970 to 2019 are shown in table 1.2. The most frequent megakillers were earthquakes (23) and hurricanes (10). Notice that 30 of the 40 worst natural disasters occurred in a belt running from China and Bangladesh through India and Iran to Turkey. Nine happened in the Americas but none were in the United States or Canada.

What is the correlation between human population density and the number of natural-disaster deaths? The data of table 1.2 paint a clear picture: densely populated Asia dominates the list of fatalities. The Asian experience offers a sobering view of what may befall the global population of humans if we continue our rapid growth. Where humans are concentrated, disasters can kill many more people during each high-energy event.

6 Chapter 1 Natural Disasters and the Human Population

TABLE 1.1 The 11 Deadliest Natural Disasters in 2018 Date **Fatalities Event** Country 2,783 28 Sep Earthquake Indonesia (7.5M)/Tsunami 483 Flood (monsoon) India (Kerala) 16 Aug 468 5 Aug Earthquake Indonesia (6.9M)430 22 Dec Volcano collapse/ Indonesia Tsunami 425 3 Jun Volcano Fuego Guatemala 220 8 Jul Floods Japan 200 31 Aug Floods Nigeria 180 Pakistan 22 May Heat wave 151 Floods North Korea 1 Sep 145 25 Feb Earthquake Papua New (7.5M)Guinea 126 23 Jul Wildfire Greece (Attica) 5,611 Total deaths

Source: Munich Reinsurance Company (2019)

THE ROLE OF GOVERNMENT IN NATURAL-DISASTER DEATH TOTALS

As the global population of humans increases, the number of deaths by natural disasters is expected to rise, but the relationship has complexities. Analyses by Gregory van der Vink and students at Princeton University show that between 1964 and 1968, about 1 person in 10,000 was killed by a natural disaster. Between 2000 and 2004, even though the population of humans doubled, the death rate by natural disaster dropped to about 1 person in 100,000. Yet, great natural disasters still result in horrific death totals in some countries. What relationships, in addition to population size, explain the locations of great natural disasters? Their study compared natural-disaster deaths to the levels of democracy and economic development within 133 nations with populations greater than 1 million that experienced five or more natural disasters between 1964 and 2004. Democracy is assessed by the World Bank's Democracy Index, and economic development by gross domestic product (GDP).

The Princeton researchers state that more than 80% of deaths by natural disasters between 1964 and 2004 took place in 15 nations, including China, Bangladesh, and Indonesia. For these 15 countries, 87% are below the median democracy index and 73% are below the median GDP. The correlation between high GDP and low death totals shows exceptions in Iran and Venezuela, two oil-rich



TABLE '	1.2		
The 40 De	eadliest Natural C	Disasters, 1970–2019	
Fatalities	Date/Start	Event	Country
300,000	14 Nov 1970	Hurricane (Bhola)	Bangladesh
255,000	28 Jul 1976	Earthquake (Tangshan)	China
245,000	26 Dec 2004	Earthquake and tsunami	Indonesia, Sri Lanka, India, Thailand
230,000	12 Jan 2010	Earthquake	Haiti
140,000	2 May 2008	Hurricane Nargis	Myanmar
140,000	29 Apr 1991	Hurricane Gorky	Bangladesh
88,000	8 Oct 2005	Earthquake	Pakistan
87,500	12 May 2008	Earthquake (Sichuan)	China
66,000	31 May 1970	Earthquake and debris flow (Nevados Huascaran)	Peru
55,630	15 Jun 2010	Heat wave and fire	Russia
50,000	21 Jun 1990	Earthquake (Gilan)	Iran
35,000	Aug 2003	Heat wave	Europe
27,000	26 Dec 2003	Earthquake (Bam)	Iran
25,000	7 Dec 1988	Earthquake	Armenia
25,000	16 Sep 1978	Earthquake (Tabas)	Iran
23,000	13 Nov 1985	Volcanic eruption and mudflows (Nevado del Ruiz)	Colombia
22,000	4 Feb 1976	Earthquake	Guatemala
20,103	26 Jan 2001	Earthquake (Gujarat)	India
19,184	11 Mar 2011	Earthquake and tsunami	Japan
19,118	17 Aug 1999	Earthquake (Izmit)	Turkey
18,000	15 Dec 1999	Flooding and debris flows	Venezuela
15,000	19 Sep 1985	Earthquake (Mexico City)	Mexico
15,000	1 Sep 1978	Floods (monsoon rains in north)	India
15,000	29 Oct 1999	Hurricane (Orissa)	India
11,000	22 Oct 1998	Hurricane Mitch	Honduras
11,000	25 May 1985	Hurricane	Bangladesh
10,800	31 Oct 1971	Hurricane (Odisha)	India
10,000	20 Nov 1977	Hurricane (Andhra Pradesh)	India
9,500	30 Sep 1993	Earthquake (Marashtra state)	India
8,960	25 Apr 2015	Earthquake	Nepal
8,135	8 Nov 2013	Hurricane Haiyan	Philippines
8,000	16 Aug 1976	Earthquake (Mindanao)	Philippines
6,425	17 Jan 1995	Earthquake (Kobe)	Japan
6,304	5 Nov 1991	Hurricane Thelma (Uring)	Philippines
6,000	Jun 1976	Heat wave	France
5,778	21 May 2006	Earthquake (Bantul)	Indonesia
5,748	14 Jun 2013	Floods	India
5,422	30 Jun 1976	Earthquake (West Irian)	Indonesia
5,374	10 Apr 1972	Earthquake (Fars)	Iran
5,300	28 Dec 1974	Earthquake	Pakistan

Source: Data from Swiss Reinsurance Company (2020).

2,059,281 Total deaths

abb91168_ch01_005-022.indd 7

Human Fatalities and Economic Losses in Natural Disasters 7



08/11/21 7:30 PM



nations with significant GDP but low democracy indices. These exceptions suggest a greater importance for democracy than GDP: the stronger the democracy index, the lower the death totals from natural disasters. The mega-killer natural disasters of recent years fit this trend also: Pakistan earthquake in 2005 (88,000 dead), Myanmar cyclone in 2008 (140,000 dead), China earthquake in 2008 (87,500 dead), and Haiti earthquake in 2010 (230,000 dead).

In a thought-provoking paragraph in their conclusion, van der Vink and students state: "Deaths from natural disasters can no longer be dismissed as random acts of nature. They are a direct and inevitable consequence of high-risk land use and the failures of government to adapt or respond to such known risks."

HUMAN RESPONSES TO DISASTER

Decades of social science research help us understand how most human beings react to natural disasters, and the news is good. Our behavior in ordinary times changes following disasters. In day-to-day life, most people are primarily concerned with their own needs and those of their immediate families; other relationships tend to be more superficial. After a natural disaster, many people change from inward-directed concerns to outward-directed actions. After an initial response of shock and disbelief, our emotions of sympathy and empathy tend to dominate. Personal priorities may be set aside and humanitarian and community-oriented actions take over. People reach out to others; they give aid and comfort to strangers; they make great efforts to provide help. Following a natural disaster, people become better connected and cohesive; they experience a heightened and compelling desire to add to the common good.

ECONOMIC LOSSES FROM NATURAL DISASTERS

The deaths and injuries caused by natural disasters grab our attention and squeeze our emotions, but in addition, there are economic losses. The destruction and disabling of buildings, bridges, roads, power-generation plants, and transmission systems for electricity, natural gas, and water, plus all the other built works of our societies, add up to a huge dollar cost. But the economic losses are greater than just damaged structures; industries and businesses are knocked out of operation, causing losses in productivity and wages for employees left without places to work.

Insured Portion of Economic Losses

Insurance companies' coverage of natural disasters extends to one-third to one-half of economic losses (table 1.3). The greatest economic losses for individual natural disasters come mostly from hurricanes and secondarily from

TABLE 1.3Insured Portion of Economic Losses (in billions of U.S. dollars)

Year	Economic Losses	Insured Portion
2019	150	52
2018	160	80
2017	340	138

Source: Munich Reinsurance Company (2020)

earthquakes. The total insurance payments for natural disasters are mostly due to severe weather via summer hurricanes and winter storms. The highest insurance dollar payouts occur in the United States, Europe, and Japan. Wealthy countries are better insured.

The locations of the worst dollar losses are different from the worst locations for fatalities (see table 1.2). Natural disaster deaths are concentrated in densely populated Asia

Natural Hazards

Many sites on Earth have not had a natural disaster in recent time, but are hazardous nonetheless. **Natural hazards** may be assessed as the probability of a dangerous event occurring. For example, people migrate and build next to rivers that are likely to flood, on the shoreline of the sea awaiting a powerful storm, and on the slopes of volcanoes that will eventually erupt. Decades, or even centuries, may pass with no great disasters, but the hazard remains.

Sites with natural hazards must be studied and understood. Their risks must be evaluated. Then we can try to prevent natural hazards from causing natural disasters. Remember: *Natural hazards are inevitable, but natural disasters are not.*

In the process of **mitigation**, we make plans and take actions to eliminate or reduce the threat of future death and destruction when natural hazards suddenly become great threats. The mitigating actions taken to protect us may be engineering, physical, social, or political.

Another need for mitigation occurs after great disasters, because people around the world tend to reoccupy the same site after a disastrous event is done. Earthquakes knock cities down, and then the survivors may use the same bricks and stones to rebuild on the same site. Floods and hurricanes inundate towns, but people return to refurbish and again inhabit the same buildings. Volcanic eruptions pour huge volumes of magma and rock debris onto the land, burying cities and killing thousands of people, yet survivors and new arrivals build new towns and cities on top of their buried ancestors. Why do people return to a devastated

8 Chapter 1 Natural Disasters and the Human Population



site and rebuild? What are their thoughts and plans for the future? For a case history of a natural hazard, let's visit Popocatépetl in Mexico.

POPOCATÉPETL VOLCANO, MEXICO

Popocatépetl is a 5,452 m (17,883 ft) high volcano that lies between the huge populations of Mexico City (largest city in Mexico) and Puebla (fourth largest city in Mexico) (figure 1.1). The volcano has had numerous small eruptions over thousands of years; thus its Nahuatl name, Popocatépetl, or Popo as it is affectionately called, means smoking mountain. But sometimes Popo blasts forth with huge eruptions that destroy cities and alter the course of civilizations. Around the year 822 CE (common era), Popo's large eruptions buried significant cities. Even its smaller eruptions have affected the course of human affairs. In 1519, Popo was in an eruptive sequence as Hernán Cortéz and about 500 Spanish conquistadors marched westward toward Tenochtitlan, the Aztec capital city. The superstitious Aztec priestking Montezuma interpreted the eruptions as omens, and they affected his thinking on how to deal with the invasion.

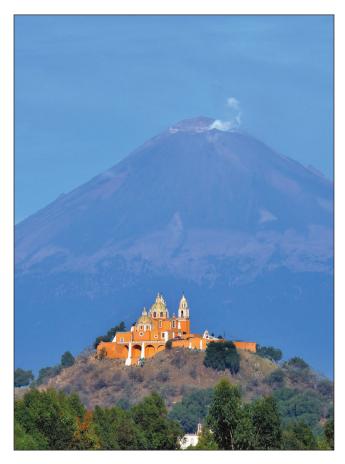


Figure 1.1 Popocatépetl in minor eruption. The cathedral was built by the Spanish on top of the great pyramid at Cholula, an important religious site in a large city that was mostly buried by an eruption around 822 CE.

Florian Kopp/imageBROKER/Alamy Stock Photo

Popocatépetl has helped change the path of history, but what is the situation now? Today, about 100,000 people live at the base of the volcano; they have been attracted by the rich volcanic soil, lots of sunshine, and fairly reliable rains. Millions more people live in the danger zone extending 40 km (25 mi) away.

Volcanic activity on Popo resumed on 9 January 2005 with eruptions of ash and gases. The sequence of intermittent eruptions continues today. How do we evaluate this hazard? Is this just one of the common multiyear sequences of small eruptions that gave the volcano its name? Or might these little eruptions be forewarnings of a giant killing eruption that will soon blast forth? We cannot answer these questions for sure. How would you handle the situation? Would you order the evacuation of 100,000 people to protect them, and in so doing, have them abandon their homes, sell their livestock, and leave their independent way of life for an unknown length of time that could be several years? Or would you explain the consequences of an unlikely but possible large eruption and let them decide whether to stay or go? If they decide to stay and then die during a huge volcanic blast, would this be your fault?

It is relatively easy to identify natural hazards, but as the Popocatépetl case history shows, it is not easy to decide how to answer the questions presented by this volcanic hazard. We are faced with the same types of questions again and again for earthquakes, landslides, tornadoes, hurricanes, floods, and wildfires.

MAGNITUDE, FREQUENCY, AND RETURN PERIOD

Earth is not a quiet and stable body. Our planet is dynamic, with major flows of energy. Every day, Earth experiences earthquakes, volcanic eruptions, landslides, storms, floods, wildfires, meteorite impacts, and extinctions. These energy-fueled events are common, but their **magnitudes** vary markedly over space and time.

Natural hazards and disasters are not spaced evenly about Earth. Some areas experience gigantic earthquakes and some areas are hit by powerful hurricanes; some are hit by both, while other areas receive neither.

During a period of several years or even several decades, a given area may experience no natural disasters. But given enough time, powerful, high-energy events will occur in every area. It is the concentrated pulses of energy that concern us here, for they are the cause of natural disasters—but how frequent are the big ones? In general, there is an inverse correlation between the **frequency** and the magnitude of a process. The frequent occurrences are low in magnitude, involving little energy in each event. As the magnitude of an event increases, its frequency of occurrence decreases. For all hazards, small-scale activity is common, but big events are rarer. For example, clouds and rain are common, hurricanes are uncommon.

Natural Hazards



Another way of understanding how frequently the truly large events occur is to match a given magnitude event with its **return period**, or recurrence interval, which is the number of years between same-sized events. In general, the larger and more energetic the event, the longer the return period.

A U.S. Geological Survey mathematical analysis of natural-disaster fatalities in the United States assesses the likeliness of killer events. Table 1.4 shows the probabilities of 10- and 1,000-fatality events for earthquakes, hurricanes, floods, and tornadoes for 1-, 10-, and 20-year intervals, and estimates the return times for these killer events. On a yearly basis, most low-fatality events are due to floods and tornadoes, and their return times are brief, less than one year. High-fatality events are dominantly hurricanes and earthquakes, and their return times for mega-killer events are much longer than for floods and tornadoes.

Knowing the magnitude, frequency, and return period for a given event in a given area provides useful information, but it does not answer all our questions. There are still the cost-benefit ratios of economics to consider. For example, given an area with a natural hazard that puts forth a dangerous pulse of energy with a return period of about 600 years, how much money should you spend constructing a building that will be used about 50 years before being torn down and replaced? Will your building be affected by a once-in-600-year disastrous event during its 50 years? Should you spend the added money necessary to guarantee

TABLE 1.4 Probability Estimates for 10- and 1,000-Death Natural Disasters in the United States						
Likeliness of a 10-Fatality Event						
	During 1 Year	During 10 Years	During 20 Years	Return Time (in years)		
Earthquake	11%	67%	89%	9		
Hurricane	39	99	>99	2		
Flood	86	>99	>99	0.5		
Tornado	96	>99	>99	0.3		
	Likeliness of a 1,000-Fatality Event					
	During 1 Year	During 10 Years	During 20 Years	Return Time (in years)		
Earthquake	1%	14%	26%	67		
Hurricane	6	46	71	16		
Flood	0.4	4	8	250		

Source: US Geological Survey Fact Sheet (unnumbered)

0.6

Tornado

that your building will withstand the rare destructive event? Or do economic considerations suggest that your building be constructed to the same standards as similar buildings in nearby nonhazardous areas?

ROLE OF POPULATION GROWTH

The world experiences significant numbers of great natural disasters and increasing economic losses from these events. The losses of life and dollars are occurring at the same time the global population of humans is increasing (figure 1.2). Population growth places increasing numbers of people in hazardous settings. They live and farm on the slopes of active volcanoes, build homes and industries in the lowlands of river floodplains, and move to hurricane-prone coast-lines. How have the numbers of people grown so large? The present situation can best be appreciated by examining the record of population history.

Overview of Human Population

The most difficult part of human history to assess is the beginning, because there are no historic documents and the fossil record is scanty. Fossils of modern humans (*Homo sapiens*) discovered in Ethiopia and South Africa are dated older than 160,000 years. Our species appears to have begun in Africa about 200,000 years ago. The rate of population growth and the number of people alive early in human history were so small that they cannot be plotted accurately on the scale of figure 1.3. The growth from a few thousand people 160,000 years ago to more than 7.8 billion people in 2020 did not occur in a steadily increasing, linear fashion. The growth rate is exponential. In its beginning, exponential growth seems negligible, then suddenly it becomes overwhelming.



Figure 1.2 The number of people on Earth continues to grow rapidly.

10 Chapter 1 Natural Disasters and the Human Population

Pat Abbott

167



THE POWER OF AN EXPONENT ON GROWTH

The most stunning aspect of figure 1.3 is the peculiar shape of the human population curve; it is nearly flat for most of human time and then abruptly becomes nearly vertical. The marked upswing in the curve shows the result of **exponential growth** of the human population. Possibly the least appreciated concept of present times is what a growth-rate exponent does to the size of a population over time. Exponential growth moves continuously in ever-increasing increments; it leads to shockingly large numbers in surprisingly short times. Probably our most familiar example of exponential growth occurs when interest is paid on money.

It can be difficult to visualize the results of exponential growth when it is expressed only as a percentage over time, such as the very small growth rate of the human population in 160,000 years or as 7% interest on your money for 50 years. It is easier to think of exponential growth in terms of doubling time—the number of years required for a population to double in size given an annual percentage growth rate. A simple formula, commonly called the rule of 70, allows approximation of doubling times:

Doubling time (in years) =
$$\frac{70}{\% \text{ growth rate/year}}$$

Learning to visualize annual percentage growth rates in doubling times is useful whether you are growing your money in investments or wasting it by paying interest on debts (especially at the high rates found with credit-card debt). Table 1.5 shows how interest rates affect how quickly your money will grow.

TABLE 1.5				
Doubling Times at Some Common Percentage Rates				
Growth Rate (% per year)	Doubling Time (years)			
0.02	3,500			
0.5	140			
1	70			
1.2	58			
2	35			
5	14			
7	10			
10	7			
17	4			

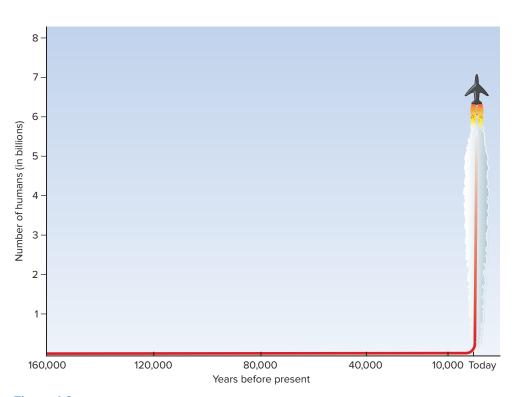


Figure 1.3 Human population growth during the past 160,000 years.

08/11/21 7:30 PM



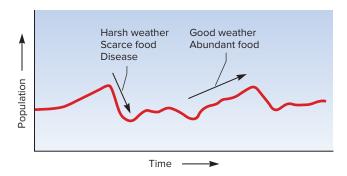


Figure 1.4 Good weather and plentiful food cause upsurges in population; bad weather, disease, and scarce food cause downswings in population.

THE PAST 10,000 YEARS OF HUMAN HISTORY

The long, nearly flat portion of the population curve in figure 1.3 certainly masks a number of small-scale trends, both upward and downward. The fossil record is not rich enough to plot a detailed record, but surely at times when weather was pleasant and food from plants and animals was abundant, the human population must have risen (figure 1.4). Conversely, when weather was harsh, food was scarce, and diseases were rampant, the human population must have fallen.

The nearly flat population growth curve began to rise about 8,000 years ago, when agriculture became established and numerous species of animals were domesticated. The world population is estimated to have been about 8 million people by 10,000 years ago. After the development of agriculture and the taming of animals removed much of the hardship from human existence, the population growth rate is likely to have increased to about 0.036% per year, yielding a net gain of 360 people per million per year. This increased rate of population growth probably caused the human population to reach 200 million people by 2,000 years ago.

As humans continued to improve their ability to modify the environment with better shelter and more reliable food and water supplies, the world population grew at faster rates. From about 1 CE to 1750, world population grew to about 800 million. Growth occurred at an average rate of 0.056% per year, meaning that another 560 people were added per million per year.

Throughout the history of the human race, high rates of birth were required to offset high rates of infant mortality and thus maintain a viable-sized human population. The 18th century saw many of the intellectual advances that set the stage for the present phase of cultural change. At long last, the causes of many diseases were being recognized.

The health necessities of clean water, sanitation, and nutrition led to the principles of public health being established. Advances in the medical world, including immunization, greatly improved the odds for the survival of individual humans through their reproductive years. No longer were many mothers and great numbers of children dying during childbirth and infancy.

The 18th century saw death rates drop dramatically, but birth rates remained high and population doubling times dropped dramatically; thus population size soared. About 1804, the human population reached 1 billion; by 1922, it had grown to 2 billion; in 1959, it reached 3 billion; by 1974, it was 4 billion; by early 1987, it was 5 billion; in 1999, it reached 6 billion; it passed 7 billion in October 2011 (figure 1.5). Notice the decline in the number of years it takes for a net gain of another 1 billion people on Earth.

Since 1900, more years have been added to average human life spans than ever before. First, the increase came by reducing child mortality. Second, the number of years in old age increased. Before 1800, no country on Earth had an average human life span at birth that exceeded 40 years. Now the average life span in every country is greater than 40 years.

The 20th-century growth of the human population is unprecedented and breathtaking. The number of humans doubled twice—from about 1.5 billion to 3 billion and again to more than 6 billion. The increased population used 16 times more energy, increased industrial output 40 times, used 7 times more water, caught 35 times more fish, and expanded the cattle population to 1.4 billion. The effect of exponential growth is racing ahead. In his book *Wealth of Nations*, published in 1776, Adam Smith said, "Men, like all other animals, naturally multiply in proportion to the means of their subsistence."

THE HUMAN POPULATION TODAY

At present, the world population is growing at about 1.1% per year for a doubling time of 63 years (table 1.6). The 1.1% gain is a net figure derived by measuring the birth rate (**fertility** rate) and subtracting the death rate (**mortality** rate). Even after subtracting all the human lives lost each year to accidents, diseases, wars, and epidemics such as AIDS and Covid-19, the human population still grows by more than 80 million people per year. Each year, the world population increases by about the total population of Germany.

The net growth of the human population can be grasped by viewing it on short timescales (figure 1.6). There is a net addition of 2.6 people every second, a rate comparable to a full jetliner landing a load of new people every minute. The monthly net growth of people is about the size of the population of Massachusetts.



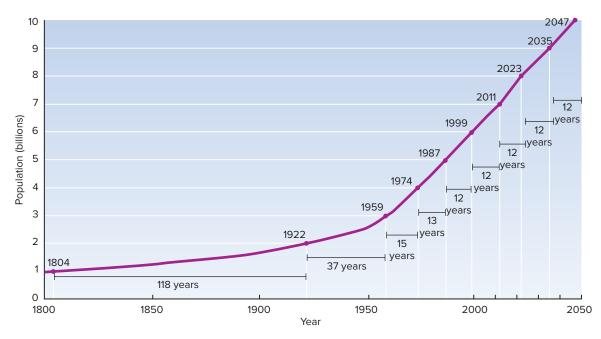


Figure 1.5 Growth of the world population of humans. Notice how the time to add another billion people has decreased. Population grew to 7.8 billion in 2020 but will exceed 10 billion before 2050.

Source: US Census Bureau.

TABLE 1.6						
World Population Data, Mid-2020						
	Population (millions)	Birth Rate (per 1,000)	Death Rate (per 1,000)	Yearly Growth %	Doubling Time (in years)	Projected Population in 2050 (millions)
World	7,773	19	7	1.1	63	9,876
More-developed countries	1,272	10	10	0	_	1,317
Less-developed countries	6,501	20	7	1.4	50	8,559
Least-developed countries*	1,062	33	7	2.6	27	1,975
Africa	1,338	34	8	2.6	27	2,560
Asia	4,626	17	7	1.0	70	5,331
Europe	747	10	11	-0.1	_	729
Northern America	368	11	8	0.3	233	435
Latin America	651	16	6	1.0	70	759
Oceania	43	17	7	1.0	70	63

^{*}Subset of less-developed countries

Source: World Population Data Sheet (2020).

Future World Population

Today, most of the more-developed countries have gone through **demographic transitions**; they have gone from high death rates and high birth rates to low death rates and low birth rates. But many less-developed countries have low

to moderate death rates and high birth rates; will they go through demographic transitions? In demographic transition theory, both mortality and fertility decline from high to low levels because of economic and social development. Yet even without significant economic development, Population Reference Bureau estimates of the rates of world

Future World Population

13





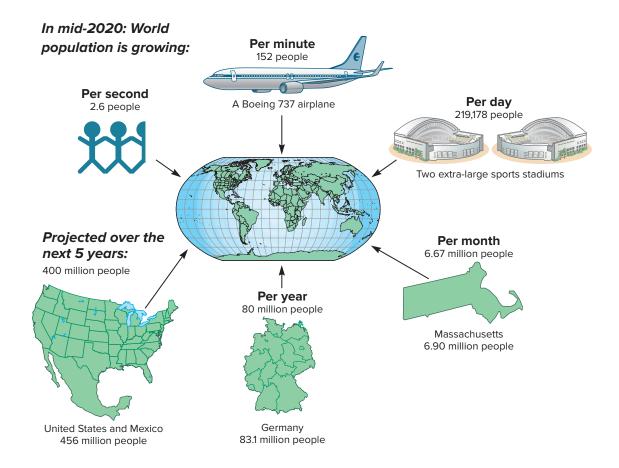


Figure 1.6 Growth of world population over differing lengths of time. Source: Modified from US Census Bureau.

population growth are dropping: from 1.8% in 1990, to 1.6% in 1997, to 1.4% in 2000, and to 1.1% in 2020. What is causing this decrease in fertility? It appears to be due largely to urbanization and increased opportunities for women. At the beginning of the 20th century, less than 5% of people in less-developed countries lived in cities, but by the year 2020, more than half of the people were living in urban areas (table 1.7). This is a change from farmer parents wanting many children to work in the fields and create surplus food, to city parents wanting fewer children to feed, clothe, and educate. Urban women have greater access to education, health care, higher incomes, and family-planning materials. When presented with choices, many women choose to have fewer children and to bear them later. Both of these choices lower the rate of population growth.

In the last 50 years of the 20th century, population grew from about 2.5 billion to over 6 billion, an increase of 3.5 billion people. Even with the recent decreases in fertility rates, the population explosion is not over. A growth rate of 1.1% per year will cause the world population of humans to exceed 10 billion by the year 2050 (see table 1.6), an increase of another 3.5 billion people within 50 years. Population growth is not evenly distributed around the world. In general, wealthy countries have low or even negative rates

of population growth. Many poor nations have high rates of population growth (figure 1.8).

An important factor in estimating future growth is the age distribution of the population (table 1.7). Nearly 30% of the population today is less than 15 years old, meaning their prime years for childbearing lie ahead. The century from 1950 to 2050 will see the world population grow from 2.5 billion to more than 10 billion people.

The number of births per woman has a dramatic effect on human population growth. Starting in the year 2000 with a world population in excess of 6 billion people, look at three scenarios for population size in the year 2150 based on births per woman: (1) if women average 1.6 children, world population drops to 3.6 billion; (2) if women average 2 children, population grows to 10.8 billion; (3) if women average 2.6 children, population grows to 27 billion. The difference between a world population of 3.6 billion or 27 billion rests on a difference of only one child per woman.

DEMOGRAPHIC TRANSITION

The demographic transition model is based on the population experiences of economically wealthy countries in the past few centuries. Up through the 17th century, a woman

14 Chapter 1 Natural Disasters and the Human Population



Side Note

Interest Paid on Money: An Example of Exponential Growth

Compare the growth of money in different situations (figure 1.7). If \$1,000 is stashed away and another \$100 is added to it each year, a linear growth process is in operation. Many of the processes around us can be described as linear, such as the growth of our hair or fingernails.

If, in contrast, another \$1,000 is stashed away but this time earns interest at 7% per year and the interest is allowed to accumulate, then an exponential growth-rate condition exists. Not only does the \$1,000 earn interest, but the interest from prior years remains to earn its own interest in compound fashion.

Notice that an exponential growth curve has a pronounced upswing, or J shape. A comparison of the linear and exponential curves in figure 1.7 shows that they are fairly similar in their early years, but as time goes on, they become remarkably different. The personal lesson here is to invest money now. Smaller amounts of money invested during one's youth will become far more important than larger amounts of money invested later in life. Individuals who are disciplined enough to delay some gratification and invest money while they are young will be wealthy in their later years. Albert Einstein described compound interest, the exponential growth of money, as one of the most powerful forces in the world.

Here is a riddle that illustrates the incredible rate of exponential growth; it shows the significance of doubling times in the later stages of a system. Suppose you own a pond and add a beautiful water lily plant that doubles in size each day. If the lily is allowed to grow unchecked, it will cover the pond in 30 days and choke out all other life-forms. During the first several days, the lily plant seems small, so you decide not to worry about cutting it back until it covers half the pond. On what day will that be?



Figure 1.7 Amounts of money versus time. Compound interest (exponential growth) produces truly remarkable sums if given enough time.

TABLE 1.7 **Data Influencing Future Population, Mid-2020** Percent of **Average Number Percent** Percent of **Population of Age** of Children Born **Urban** (cities **Married Women Using** per Woman >2,000 people) **Modern Contraception** <15 World 26 2.3 56 55 19 79 More-developed 16 1.6 61 countries Less-developed 28 7 54 2.5 51 countries Least-developed 40 4.1 34 34 countries 32 Africa 41 3 4.4 43 Asia 24 9 2.0 51 57 19 75 1.5 63 Europe 16 Northern America 18 17 82 68 1.7 24 9 2.0 79 71 Latin America 12 23 2.3 68 56

Source: World Population Data Sheet (2020).

Oceania



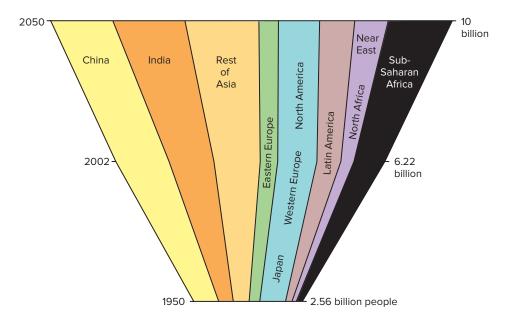


Figure 1.8 World population by region: 1950, 2002, 2050.

Source: US Census Bureau

had to bear several children to have a few survive to adulthood and replace the prior generation. Births had to be numerous to compensate for the high rates of infant mortality. Beginning in the 18th century, discoveries in public health, medicine, and immunization caused the death rate to drop dramatically. During this time, birth rates stayed high, so overall population grew rapidly. As time passed and people realized that most of their children would survive to adulthood, birth rates dropped and population stabilized at a new and higher level.

The demographic transition takes place in phases:

- 1. Before the transition: high death rates are offset by high birth rates to maintain a population.
- During the transition: low death rates coupled with continuing high birth rates cause population to soar.
- After the transition: low death rates combine with low birth rates to achieve a stable population at a significantly higher level.

Today the transition is taking place at different rates in different countries (figure 1.9). Most of the population growth is occurring in the poorest areas of the poorest countries. Some of the wealthiest countries now have more deaths than births each year.

URBANIZATION AND EARTHQUAKE FATALITIES

During the past 500 years, global earthquakes killed about 5 million people. Average numbers of deaths were about 1 million per century, or 100,000 per decade. These simple averages are misleading because they hide the effects of the

Demographic transition

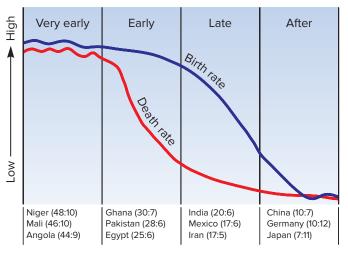


Figure 1.9 Demographic transition. In mid-2020 the shifts in birth rates and death rates vary markedly between countries. Birth and death rates are both expressed in number of people per 1,000 each year. For example, Niger has 48 births and 10 deaths per 1,000 people each year (48:10).

Data from Population Reference Bureau.

deadliest earthquakes, such as the 250,000 people killed by the Tangshan, China, event in 1976. The mega-killer earthquakes of the past 500 years occurred in China, Indonesia, Pakistan, Iran, Turkey, Italy, Japan, and Haiti—and they may occur there again.

An analysis of the past 500 years by Roger Bilham shows that, with an average population of about 1.5 billion people, there was one earthquake that killed nearly a million people.

16 Chapter 1 Natural Disasters and the Human Population



A Classic Disaster

Influenza (Flu) Pandemic of 1918

In July 1914, war broke out in Europe and grew into World War I. The countries and empires involved contained about half the people in the world. When the war ceased in November 1918, about 9 million soldiers and 7 million citizens were dead as a direct result of the fighting.

As bad as 16 million war deaths sound, a far more deadly natural disaster began during that time—the influenza pandemic of 1918–1919. The flu pandemic killed more than 50 million people; this was 3% of the world's population. Estimates of total deaths range up to 100 million people. The influenza migrated around the world in waves. The first wave began in the spring of 1918 and was relatively mild; it caused few deaths. But viruses are not static; they mutate and change. The lethal second wave arrived in the fall of 1918 and swept through the world; about two-thirds of the deaths worldwide occurred between late September and late December. Victims' lungs fill with fluids, causing death by suffocation. A third wave occurred in

March and April 1919. Most of the victims were healthy young adults between 18 and 45 years of age—rather than typical flu victims, who are seniors or juveniles. The strategies used to fight the flu in 1918 sound familiar today—isolation, quarantine, good personal hygiene, use of disinfectants, limited public gatherings, and wear masks.

In 1918 the U.S. population was about 103 million. Estimates are that about 25% of those people caught the flu, and about 675,000 of them died. This is a kill rate of 2.7%.

World War I did not cause the flu, but the global movements of millions of troops, weakened by stress and battle, increased the spread and deadly effects of the **virus**. Another 3 million soldiers died, not from World War I battles, but from influenza. In 1918, children skipped rope to this rhyme:

I had a little bird Its name was Enza I opened the window And in-flu-enza.

But with population becoming five times larger at 7.8 billion people in the year 2020, million-death earthquakes may occur five times as frequently, or about one per century. Most of the human population growth, by birth and by migration, is occurring in cities in less-developed countries. Many of these people are living in poorly constructed buildings in mega-cities. Million-death earthquakes are possible in a growing number of mega-cities.

DISEASE PANDEMICS

Throughout recorded history, deadly diseases have swept throughout the world, killing millions of people in **pandemics.** For example, the bacterium *Yersinia pestis*, transmitted to humans by fleas, caused the bubonic plague—the Black Death that killed about 75 million people in Europe in the 14th century.

Viruses have also caused pandemics via smallpox, HIV, polio, **influenza**, and other diseases. For example, in 1918–1919, with population less than 2 billion, the influenza virus A spread around the world, killing more than 50 million people. Now, with the human population about 8 billion people, with more than 50% of people living in cities, and with the rapid movement of people worldwide via jet airplanes, the potential for new pandemic diseases is high.

Viruses

Viruses are life in the simplest form. They have no cells, no bodies, no metabolism of their own. They are genetic material (DNA or RNA) coated by fat and protein. A **virus** might have only 4 genes, whereas a bacterium might have 4,000 genes, and a human 24,000 genes. Viruses cannot reproduce by themselves; they must invade a host cell and

cause the host to reproduce the virus. A virus is alive only in the cells of others; it is a virtual organism (software) running on borrowed hardware (for example, you) to produce more copies of its genome. Viruses have killed more people, killed more living organisms, than any other type of predator.

Zoonotic Diseases Viruses infect many forms of life, including animals, plants, and bacteria. We humans share a common pool of viruses with other organisms. Disease-causing viruses transmitted into humans from other species give us **zoonotic diseases**. For example, the same viruses are commonly found in humans, pigs, and birds. There are an estimated 1 billion pigs and 20 billion chickens alive today. Because humans commonly live and interact with pigs and birds, the three-way transfer of viruses among us occurs frequently.

Most new human diseases each year are zoonotic; they are **spillover events** from non-human animal hosts. Viruses can come directly from animals, domestic or wild, or they can be carried by an intermediate species (a *vector*) that is not sick. Viruses transferred to humans include HIV/AIDS from chimpanzees, Ebola from bats, SARS from bats through civets (mongoose), and MERS through camels. After different viruses enter a single cell, their genes can form new combinations, creating new types of viruses.

As the rapidly growing human population expands into wildlands, there are increasing chances for **zoonosis**. Farms on the edge of forests, duck farms near lakes, markets where wild animals are sold to be eaten—these are places where zoonotic diseases are likely to emerge. For example, in 1999, fruit bats living in a Malaysian forest dropped partially eaten fruit into a pig farm next to the forest. Pigs ate the bat saliva

Future World Population







covered fruit, and then the pigs transferred the deadly Nipah virus into humans.

Bats are unique creatures; they are the only flying mammal, and they have an amazing ability to host viruses without becoming diseased. Bats have evolved an internal system using both defense to, and tolerance of, viruses; they are the reservoir host of innumerable viruses. There are 1,422 recognized species of bats; they exist in all environments except polar regions. Bats provide valuable services to life such as pollination, seed dispersal, and insect control. Humans could learn much from bats about virus control

Influenza A Viruses Influenza A viruses cause recurrent **epidemics** and pandemics, as in 1918–1919. Type A viruses examined on the basis of their haemagglutinin (HA) and neuraminidase (NA) molecules are divided into 18 HA subtypes (H1 to H18) and 11 NA subtypes (N1 to N11). In 2005, researchers reported the results of a study of the 1918–1919 influenza virus collected from samples preserved from World War I flu-victim soldiers and from historic individuals buried in Arctic permafrost (frozen soil). The 1918–1919 influenza was type A (H1N1), a subtype with an early history in birds.

Early in 2009, a flu epidemic broke out near La Gloria in the state of Veracruz, Mexico. By 23 April 2009, 23,000 cases had been reported. By 7 May 2009, the flu had spread to become a pandemic, with cases identified in 21 countries on five continents. Laboratory analyses showed that this new virus was type A (H1N1) and was made up of genes from four different flu viruses: from North American pigs (30.6%), Eurasian pigs (17.5%), North American birds (34.4%), and humans (17.5%). People were worried. Could this virus evolve into as big a killer as the one in 1918–1919?

Analysis of H1N1 deaths in 2009 from 214 countries showed 44,100 deaths—a significant total, but far less severe than in 1918. Like the 1918 influenza, though, most of the deaths occurred in young people; 73% of deaths were people 29 years old and younger. The death percentages by age groups include:

- 37% were 10 to 19 years old
- 22% were less than 9 years old
- 14% were 20 to 29 years old

In the reverse of a typical flu year, people 60 years old and older suffered only 3% of the deaths. If one views the 2009 H1N1 figures as deaths only, then it was not as bad a year as had been feared. But if one considers the number of years of life lost by the young victims, then the 2009 pandemic would be more equivalent to 250,000 deaths in a typical flu year. It was a disease of the young.

Coronaviruses Coronaviruses are not the same as influenza viruses; they belong to a completely different group of viruses. Coronavirus-caused diseases range from the

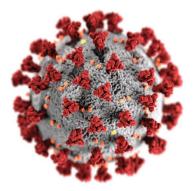


Figure 1.10 SARS-CoV-2 virus. Spikes on the outer surface of the virus link up with human cell receptors, enabling the virus to enter human cells, just like a key opening a lock.

Source: Centers for Disease Control and Prevention Alissa Eckert, MS: Dan Higgins, MAMS/CDC.

common cold to deadly respiratory infections. A disease-causing novel coronavirus was recognized in December 2019 in Wuhan, China. On 11 January 2020 the genome of the novel coronavirus was published by the Chinese. The new virus spread quickly. On 30 January 2020, after 18 countries reported infections, the World Health Organization (WHO) declared a Public Health Emergency of International Concern, its highest level of alarm. By late March 2020, in less than three months, global cases of Covid-19 reached one million. In another three months, by late June 2020, it took less than one week for a million new cases. Again, we see the effect of exponential growth. The disease spread rapidly into a global pandemic.

This novel coronavirus is SARS-CoV-2 (figure 1.10); it causes a disease called Covid-19. The new virus is 96% identical to a virus found in horseshoe bats living in Yunnan Province, China. The bat virus could evolve by simple mutations or by swapping genetic material with other bat species or within other animals such as pangolins (scaly anteaters). The city of Wuhan has wet markets where wild animals, including bats and pangolins, are sold to be eaten by people. Wet markets are incubators for viruses. Humans handling and eating wild animals may have started the pandemic.

The city of Wuhan also hosts the Wuhan Institute of Virology, a leading collector of wild bat coronaviruses. Research, such as gain-of-function experiments, could modify coronaviruses into more dangerous varieties that may have escaped from the laboratory.

Viruses can spread easily within the human population. Small incidents can have large effects on the future. As an early example, on 17 March 2020 in Mount Vernon, Washington, at a choir practice involving 61 people, one infected person unknowingly passed the virus on to 52 people; two of them died. This was a *superspreader event*.

The coronavirus initially settles in the respiratory tract—the nose, mouth, and lungs. But Covid-19 is not a typical respiratory system infection. It can affect the entire

18 Chapter 1 Natural Disasters and the Human Population



body. It commonly causes a huge inflammatory response from the immune system (a cytokine storm), which can lead to organ failure. It can invade a layer of cells inside the blood vessels, causing blood clots to form in the lungs, heart, and brain. These problems can lead to pneumonia, heart attack, stroke, organ failures, and death. It is a complicated disease that does not affect everyone in the same way. It is probably here to stay as an *endemic disease* that must be dealt with each year via vaccines and therapeutics, as we do with the flu and the common cold.

On 11 March 2021, the one-year anniversary of the declared pandemic, there were more than 118 million confirmed cases globally with more than 2.6 million deaths. The United States had 29 million confirmed cases with more than 525,000 deaths. With about 4% of the world's population the U.S. suffered more than 20% of the Covid-19 deaths.

Analysis by the U.S. Center for Disease Control in mid-October 2020 showed that U.S. Covid-19 death percentages by age groups were:

- 1% were <1 to 34 years old
- 7.4% were 35 to 54 years old
- 34.4% were 55 to 74 years old
- 57.2% were 75 and older

More than 91% of the deaths were people 55 years and older. It is a disease of the old.

Carrying Capacity

How many people can Earth support? At this time, the question is unanswerable. Nonetheless, many people worry about dangers resulting from the unprecedented growth of the human population, such as more and greater natural disasters, increased global warming, decreasing supplies of fresh water, increased pollution, increased desertification, and the increased rate of extinction of species. Other people see no big problems and point out that humans have already increased the carrying capacity of Earth via agriculture, water storage and purification, and advances in public health; they feel that any upcoming problems will be solved just like others have been in the past.

In the natural world, biologists studying **carrying capacity** of the environment for individual species of mammals, birds, frogs, and other animals find that population size is regulated by the resources available. For example, when a resource such as available food increases, a feeding population grows in size. If that food resource decreases due to drought, competition, or other causes, the population dependent on that food dies back and decreases in numbers.

A fundametal principle of biology is that a population of animals cannot increase forever because they live in a finite ecosystem. Ultimately, population growth is controlled by negative feedbacks such as starvation, predation, and disease.

Ireland in the 1840s

Ireland in the 1840s provides a human example of carrying capacity. The European explorers of the 1500s brought the potato back from South America. The potato is a highly nutritious food. A diet of potatoes, milk (from animals fed potatoes), and greens constitute a nutritionally complete diet. An acre of potatoes could feed an Irish family of six for a year. In Ireland, the potato was the wonder crop that allowed a child-loving population to grow explosively. By 1841, Ireland's population had grown well past 8 million, with nearly half the people surviving wholely or mostly on potatoes. In 1845, heavy spring rains aided growth and spread of a fungal infestation, the potato blight, which caused potatoes to rot during storage. But when the potatoes rotted there was no substitute food. Malnourishment became common. Then the winter of 1846-1847 hit with unusual severity, causing weakened people to suffer even more. The toll was severe: a million people died from disease, and another 1.5 million people emigrated. During their emigration to the United States and Canada, 1 in 7 people died.

The carrying capacity of Irish land increased for humans when the potato arrived. Potato plants covered the lands, even extending into bogs and up steep mountain slopes. The human population fed by the increased food supply grew rapidly. But when the potato supply dropped suddenly, so did the human population. The human population in Ireland today is less than it was in 1845.

Easter Island (Rapa Nui)

Easter Island (Rapa Nui) is a triangular-shaped, volcanic land with an area of about 165 km² (64 mi²). It lies over 2,000 km (1,240 mi) east of Pitcairn Island and over 3,700 km (2,300 mi) west of Chile (figure 1.11). Easter Island is isolated; it has high temperatures and humidity, poorly drained and marginal soils, no permanent streams, no terrestrial mammals, about 30 native plant species including trees in locally dense growths, and few varieties of fish in the surrounding sea. Year-round water is available only in little lakes within the volcano caldera.

About 1,000 years ago, seafaring Polynesian people arrived on Rapa Nui with 25 to 50 settlers. They were part of the great Polynesian expansion outward from southeast Asia that led them to discover and inhabit islands from Hawaii in the north to New Zealand in the southwest and to Rapa Nui in the southeast. The wide-ranging voyagers colonized islands over a Pacific Ocean area more than twice the size of the United States.

The colonizers of Rapa Nui brought chickens and rats, along with several of their food plants. The climate was too severe for most of their plants except the yam. Their

Carrying Capacity

08/11/21 7:31 PM



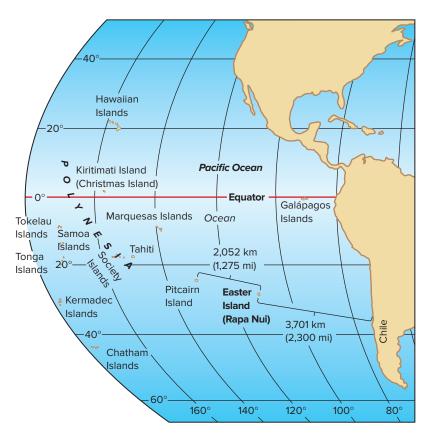


Figure 1.11 Easter Island (Rapa Nui) is an isolated outpost of Polynesian civilization nearly lost in the vast Pacific Ocean.



Figure 1.12 Rapa Nui inhabitants spent much of their energy creating giant statues (moai).

Adalberto Rios Szalay/Sexto Sol/Getty Images

resulting diet was based on easily grown chickens and yams, along with seabirds and their eggs. Housing was fashioned using wood from native trees; the people had lots of free time.

The islanders used their free time to develop a complex social system divided into clans that practiced elaborate rituals and ceremonies. Their customs included competition between the clans in shaping and erecting mammoth statues. The statues were carved out of volcanic rock using obsidian (volcanic glass) tools. The statues (moai) were more than 6 m (20 ft) high, weighed about 15 tons apiece, and were erected on ceremonial platforms (ahu) (figure 1.12).

The peak of the civilization occurred about 1550 CE, when the human population had risen to about 7,000; statues numbered more than 600, with half as many more being shaped in the quarries. But from its peak, the civilization apparently declined rapidly as first witnessed by the crew of a Dutch ship on Easter Sunday, 5 April 1722. The Europeans reported finding about 2,000 people living in caves in a primitive society engaged in warfare. What caused this cultural collapse? It appears that human activities so overwhelmed the environment that it was no longer able to support the greatly enlarged human population. The customs of society dissolved in the fight of individuals and clans to survive.

Carving the giant statues had not been particularly difficult, but transporting them was physically and environmentally strenuous. Trees were cut down and placed under statues as rollers. Islanders pushed the heavy statues from the quarry and levered them onto their ceremonial platforms. The competition between clans to create the most statues helped destroy the forests. Without trees, houses could not be built, and people had to move to caves. There was little fuel for cooking or to ward off the chill of colder times. Soil erosion increased and agricultural production dropped. Without trees, there were no canoes, and so islanders caught fewer fish. Without canoes, there was no escape from the remote and isolated island. Seabirds were overkilled, thus reducing egg collection for food and guano mined for fertilizer. As food resources declined, the social system collapsed, and the statue-based religion disintegrated. Clans were reduced to warfare and cannibalism in the struggle for food and survival.

Competition between clans was so consuming that they did not consider the health of the environment and thus paid a price: the human population on the island collapsed. Easter Island is one of the most remote inhabited areas on Earth, a tiny island virtually lost in the vast Pacific Ocean. When problems set in faster than the Rapa Nui customs could solve them, there was no place to turn for help, no place to escape.

20 Chapter 1 Natural Disasters and the Human Population



The carrying capacity of the land had been exceeded, and the human population suffered. What lesson does Easter Island have for the whole world? Earth is but a tiny island lost in the vast ocean of the universe (figure 1.13); there is no realistic chance of the human population escaping to another hospitable planet.

The Easter Island example raises interesting philosophical questions. If climate change decreases global food production, causing the human population to exceed Earth's carrying capacity, could human value systems change fast enough to solve the problem? If all the people on Earth had to face the Easter Island situation, how would we fare?



Figure 1.13 Earth is an isolated island nearly lost in the vast "ocean" of the universe.

Pat Abbott

Summary

Great natural disasters killed almost 500,000 people in four recent events: 2011 Japan earthquake and tsunami, 2010 Haiti earthquake, 2008 hurricane in Myanmar, and 2008 earthquake in China. Over time, the two deadliest events are tropical storms (hurricanes) and earthquakes. In 2019, the known economic losses from natural disasters were about US\$150 billion. The long-term trend is for economic losses to increase.

Natural hazards exist in areas of obvious danger, such as cities built on the slopes of active volcanoes or on the flood-plains of rivers. For these sites, it is only a matter of time before the hazard is realized as a disaster. At any one site, the greater the magnitude of a disaster, the less frequently it occurs. Large disasters have longer return periods.

The curve describing the history of human population growth is flat to gently inclined for 160,000 years, and then it rises rapidly in the last three centuries. In the past, women bore numerous children, but many died, so overall population growth was slow. With the arrival of the scientific-medical revolution and the implementation of the principles of public health, the human population has soared. Birth rates remain high in much of the world, even though death rates have plummeted. The population reached 1 billion in about 1804, 2 billion in 1922, 3 billion in 1959, 4 billion in 1974, 5 billion in 1987, 6 billion in 1999, and it passed 7 billion in 2011.

A steeply rising growth curve is exponential; in terms of population, more people beget ever more people. One way to visualize exponential growth is by using doubling time, the length of time needed for a population to double in size. Doubling times can be approximated by the rule of 70:

Doubling time (in years) =
$$\frac{70}{\% \text{ growth rate/year}}$$

At present, after subtracting deaths from births, world population increases 1.1% per year for a doubling time of 63 years.

Much hope is placed in the demographic transition model, which holds that economic wealth, combined with knowing that one's children will survive, leads to dramatic drops in birth rates. This model holds for some more-developed countries. Some less-developed countries are experiencing drops in birth rates, presumably due to urbanization and more choices for women. Even at lower rates of growth, human population is likely to exceed 10 billion before 2050. The rapid growth in human population sets the stage for mega-death earthquakes and hurricanes.

New flu viruses are commonly created where people live closely with birds and pigs. These new viruses, which have the potential to kill millions of people, can rapidly spread around the world.

Carrying capacity is an estimate of how many individuals of a species the environment can support. How many people can Earth support? The answer is not known, but it is the subject of much debate.

Terms to Remember

carrying capacity 19
CE 9
coronavirus 18
cyclone 6
demographic transition 13
earthquake 6
energy 6
epidemic 18
exponential growth 11

Terms to Remember



pandemic 17 return period 10 spillover event 17 tsunami 6

virus 17 volcano 9 zoonosis 17 zoonotic disease 17

Questions for Review

- 1. What types of natural disasters killed the most people in the past 40 years? Where in the world are deaths from natural disasters the highest? Where in the world are insurance losses from natural disasters the highest?
- 2. What is a great natural disaster?
- 3. What is the difference between a natural disaster and a natural hazard? How do economic losses differ from insured losses?
- 4. For nations, what is the relationship between natural-disaster deaths, gross domestic product, and level of democracy?
- 5. What is the relationship between the magnitude of a given disaster and its frequency of occurrence?
- 6. Draw a curve showing the world population of humans in the past 100,000 years. Why has the curve changed shape so dramatically?
- 7. Explain the concept of exponential growth.
- 8. What is the size of the world population of humans today? Extrapolating the current growth rate, what will the population be in 100 years? In 200 years? Are these large numbers environmentally realistic?
- 9. What are the population doubling times given these annual growth rates: Africa, 2.4%; world, 1.2%?
- 10. For nations, what are demographic transitions?
- 11. How much time does it take for a flu pandemic to infect people all around the world?

- 12. What is the relationship between earthquake fatalities and cities?
- 13. Explain the concept of carrying capacity for a species. What processes might limit the numbers of a species?
- 14. How does a virus differ from other life forms?
- 15. What are zoonotic diseases? What are some domestic animals that humans commonly transmit viruses to and receive viruses from? What are some wild animals that carry viruses that infect us?

Questions for Further Thought

- Would we call a large earthquake or major volcanic eruption a natural disaster if no humans were killed or buildings destroyed?
- 2. Which single disaster could kill the most people—a virus pandemic, an earthquake, or a hurricane?
- 3. Could global building designs be made disaster-proof, thus reducing the large number of fatalities?
- 4. What is the carrying capacity of Earth for humans—that is, how many humans can Earth support? What factors are most likely to slow human population growth?
- 5. Compare the rate of change of human populations to the rate of change in religious and cultural institutions. Can religious and cultural institutions change fast enough to deal with world population growth?
- 6. Evaluate the suggestion that the overpopulation problem on Earth can be solved by colonizing other planets.
- 7. Is a nation's destiny determined by its demographics?
- 8. Is a virus alive?





Chapter 1 Natural Disasters and the Human Population

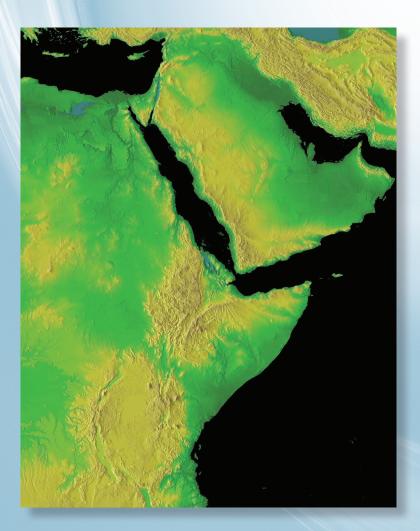


CHAPTER 2

Internal Energy and Plate Tectonics

Such superficial parts of the globe seemed to me unlikely to happen if the Earth were solid to the centre. I therefore imagined that the internal parts might be a fluid more dense, and of greater specific gravity than any of the solids we are acquainted with; which therefore might swim in or upon that fluid. Thus the surface of the globe would be a shell, capable of being broken and disordered by the violent movements of the fluid on which it rested.

—Benjamin Franklin, 1780



Satellite view of Arabia moving northeast away from Africa.

Source: NOAA/NGDC

LEARNING OUTCOMES

Internal energy has caused the Earth to differentiate into layers. Throughout the Earth, materials move vertically and horizontally. After studying this chapter you should

- · know the layering of the Earth and how it formed.
- be familiar with the sources of energy inside the Earth.
- understand the behavior of materials.
- be able to explain how plate tectonics operates.
- comprehend Earth's magnetic field and the evidence it provides for plate tectonics.
- know the age of the Earth and how it is determined.
- appreciate the thought processes used to understand the Earth.

OUTLINE

- Origin of the Sun and Planets
- Earth History
- The Layered Earth
- Internal Sources of Energy
- Age of Earth
- Plate Tectonics
- The Grand Unifying Theory
- How We Understand Earth





s described in the Prologue, energy flows upward and outward from the interior of the Earth. At the surface of the Earth we feel this energy where it is released as earthquakes and volcanic eruptions. In order to understand these natural disasters, we need to know what is going on inside the Earth and how plate tectonics operates.

Origin of the Sun and Planets

Impacts of asteroids, comets, and meteorites are not rare and insignificant events in the history of our Solar System; they probably were responsible for its formation. The most widely accepted hypothesis of the origin of the Solar System was stated by the German philosopher Immanuel Kant in 1755. He thought the Solar System had formed by the growth of the Sun and planets through collisions of matter within a rotating cloud of gas and dust.

The early stage of growth began within a rotating spherical cloud of gas, ice, dust, and other solid debris (figure 2.1a). Gravity acting upon matter within the cloud attracted particles, bringing them closer together. Small particles stuck together and grew in size, resulting in greater gravitational attraction to nearby particles and thus more collisions. As matter drew inward and the size of

the cloud decreased, the speed of rotation increased and the mass began flattening into a disk (figure 2.1b). The greatest accumulation of matter occurred in the center of the disk, building toward today's Sun (figure 2.1c). The two main constituents of the Sun are the lightweight elements hydrogen (H) and helium (He). As the central mass grew larger, its internal temperature increased to about 1,000,000 degrees **centigrade** (°C), or 1,800,000 degrees **Fahrenheit** (°F), and the process of **nuclear fusion** began. In nuclear fusion, the smaller hydrogen atoms combine (fuse) to form helium, with some mass converted to energy. We Earthlings feel this energy as **solar radiation** (sunshine).

The remaining rings of matter in the revolving Solar System formed into large bodies as particles continued colliding and fusing together to create the planets (figure 2.1d). Late-stage impacts between ever-larger objects would have been powerful enough to melt large volumes of rock, with some volatile elements escaping into space. The inner planets (Mercury, Venus, Earth, Mars) formed so close to the Sun that solar radiation drove away most of their volatile gases and easily vaporized liquids, leaving behind rocky planets. The next four planets outward (Jupiter, Saturn, Uranus, Neptune) are giant icy bodies of hydrogen, helium, and other frozen materials from the beginning of the Solar System.

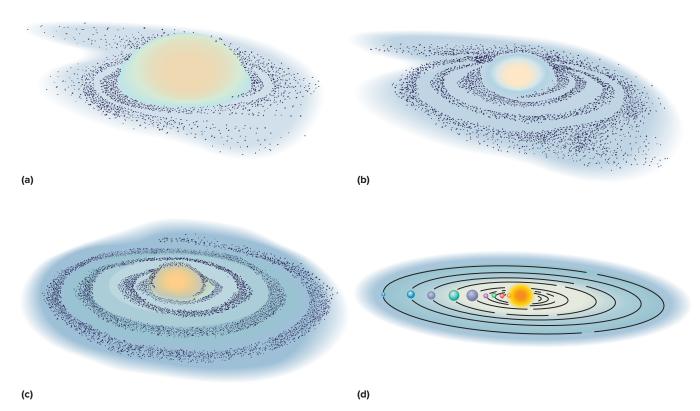


Figure 2.1 Hypothesis of the origin of the Solar System. (a) Initially, a huge, rotating spherical cloud of ice, gas, and other debris forms. (b) The spinning mass contracts into a flattened disk with most of its mass in the center. (c) Planets grow as masses collide and stick together. (d) The ignited Sun is surrounded by planets. Earth is the third planet from the Sun.



IMPACT ORIGIN OF THE MOON

Large impacts can generate enough heat to vaporize and melt rock; they can produce amazing results. For example, a long-standing hypothesis on the origin of Earth's Moon involves an early impact of the young Earth with a Mars-size body, a mass about 10 times larger than the Moon. The resultant impact generated a massive cloud of dust and vapor, part of which condensed and accumulated to form the Moon. This theory suggests that the Moon is made mostly from Earth's rocky **mantle.** The theory accounts for the lesser abundance of iron on the Moon (iron on Earth is mostly in the central **core**) and the Moon's near lack of lightweight materials (such as gases and water), which would have been lost to space.

Earth History

To understand the origin and structure of Earth, we must know the flows of energy throughout the history of our planet. Studying early history is difficult because Earth is a dynamic planet; it recycles its rocks and thus removes much of the record of its early history. The older the rocks, the more time and opportunities there have been for their destruction. Nonetheless, the remaining early Earth rocks, along with our growing knowledge of the processes in Earth's interior and in the Solar System, allow us to build an increasingly sophisticated approximation of early Earth history.

Earth appears to have begun as an aggregating mass of particles and gases from a rotating cloud about 4.6 billion years ago. During a 30- to 100-million-year period, bits and pieces of metal-rich particles (similar to iron-rich meteorites), rocks (similar to stony meteorites), and ices (composed of water, carbon dioxide, and other compounds), accumulated to form Earth. As the ball of coalescing particles

enlarged, the gravitational force may have pulled more of the metallic pieces toward the center, while some of the lighter-weight materials may have concentrated near the exterior. Nevertheless, Earth in its infancy probably grew from random collisions of debris that formed a more or less homogeneous mixture of materials.

But Earth did not remain homogeneous. The very processes of planet formation (figure 2.2) created tremendous quantities of heat, which fundamentally changed the young planet. The heat that transformed Earth came primarily from (1) impact energy, (2) decay of radioactive isotopes, (3) gravitational energy, and (4) differentiation into layers (figure 2.2).

As the internal temperature of Earth rose beyond 1,000°C (1,800°F), it passed the melting points of iron at various depths below the surface. Iron forms about one-third of Earth's mass, and although it is much denser than ordinary rock, it melts at a much lower temperature. The buildup of heat caused immense masses of iron-rich meteorites to melt. The highdensity liquid iron was pulled by gravity toward Earth's center. As these gigantic volumes of liquid iron moved inward to form Earth's core, they released a tremendous amount of gravitational energy that converted to heat and probably raised Earth's internal temperature by another 2,000°C (3,600°F). The release of this massive amount of heat would have produced widespread melting likely to have caused low-density materials to rise and form: (1) a primitive **crust** of low-density rocks at the surface of Earth; (2) large oceans; and (3) a denser atmosphere. The formation of the iron-rich core was a unique event in the history of Earth. The planet was changed from a somewhat homogeneous ball into a density-stratified mass with the denser materials in the center and progressively lessdense materials outward to the atmosphere.

The low-density materials (magmas, waters, and gases), freed by the melting, rose and accumulated on Earth's

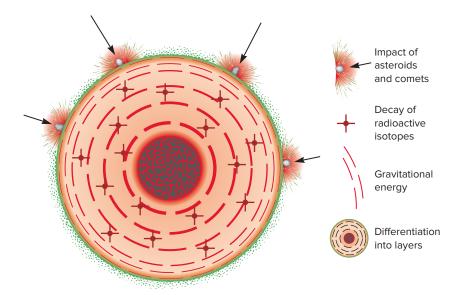


Figure 2.2 Heat-generating processes during the formative years of Earth include (1) impact energy, (2) decay of radioactive isotopes, and (3) gravitational energy. Increasing heat caused Earth to differentiate into layers.



exterior as continents, oceans, and atmosphere. It seems that oceans and small continents existed by 4.4 billion years ago, life probably was present as photosynthetic bacteria 3.5 billion years ago, large continents were present at least 2.5 billion years ago, and the outer layers of Earth were active in the process of plate tectonics by at least 1.5 billion years ago.

The Layered Earth

Earth today is differentiated into layers of varying densities. As we have noted, much of the densest material was pulled toward the center, and some of the least dense substances escaped to the surface (figure 2.3). At the center of Earth is a dense, iron-rich core measuring about 7,000 km (4,350 mi) in diameter. Earth's core is almost exactly the same size as planet Mars. The inner core is primarily a solid mass 2,450 km (1,520 mi) in diameter with temperatures of at least 4,300°C (7,770°F). The outer core is mostly **liquid**, and the **viscous** movements of convection currents within it are responsible for generating Earth's magnetic field. The entire iron-rich core is roughly analogous in composition to a melted mass of iron-rich meteorites.

Surrounding the core is a rocky mantle nearly 2,900 km (1,800 mi) thick, with a composition similar to that of stony meteorites. The mantle comprises 83% of Earth's volume and 67% of its mass. The rocks of the mantle can be approximated by melting a stony meteorite in the laboratory;

this produces a separation in which an upper froth rich in low-density elements rises above a residue of denser minerals/elements. The low-density material is similar to continental crust that by melting and separation has risen above the uppermost mantle. All the years of heat flow toward Earth's surface have "sweated out" many low-density elements to form a continental crust. Today, the continents make up only 0.1% of Earth's volume. Floating above the rocky layers of Earth are the oceans and the atmosphere.

Earth's layering can be described as based on either (1) different strengths or (2) different densities due to varying chemical and mineral compositions (figure 2.3). Both temperature and pressure increase continuously from Earth's surface to the core, yet their effects on materials are different. Increasing temperature causes rock to expand in volume and become less dense and more capable of flowing under pressure and in response to gravity. Increasing pressure causes rock to decrease in volume and become more dense and more rigid. Visualize tar at Earth's surface. On a cold day, it is solid and brittle, but on a hot day, it can flow as a viscous fluid. Similar sorts of changes in physical behavior mark different layers of Earth. In fact, from a perspective of geological disasters, the crust-mantle boundary is not as important as the boundary between the rigid **lithosphere** (from the Greek word *lithos*, meaning "rock") and the "soft plastic" asthenosphere (from the Greek word asthenes, meaning "weak") (figure 2.4). The mesosphere, the mantle below the asthenosphere (see figure 2.3), is solid; it is a "stiff plastic," but it is not brittle like the lithosphere.

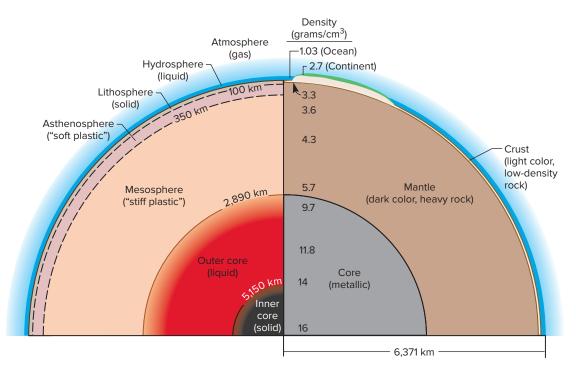


Figure 2.3 Density stratification within Earth—that is, lower-density materials float atop higher-density materials. Pressure and temperature both increase from the surface to the center of Earth. Layers illustrated on the left show the differences in physical properties and strengths. Layers on the right emphasize different mineral and chemical compositions.



Side Note

Mother Earth

The history of the 4.6-billion-year-old Earth has been metaphorically contrasted with the life history of a 46-year-old woman by Nigel Calder in his book *The Restless Earth*. In this metaphor, each of "Mother" Earth's years equals 100 million years of geologic time. The first seven of her years are mostly lost to the biographer. Like human memory, the early rock record on Earth is distorted; it emphasizes the more recent events in both number and clarity. Most of what we know of "Mother" Earth happened in the past six years of her life. Her continents had little life until she was 42. Flowering plants did not appear until her

45th year. Her pet dinosaurs died out eight months ago. In the middle of last week, some ancestors of present apes evolved into human ancestors. Yesterday, modern humans (*Homo sapiens*) evolved and began hunting other animals, and in the last hour, humans discovered agriculture and settled down. Fifteen minutes ago, Moses led his people to safety; five minutes later, Jesus was preaching along the same fault line; and after another minute, Muhammad taught in the same region. In the last minute, the Industrial Revolution began, and the number of humans increased enormously.

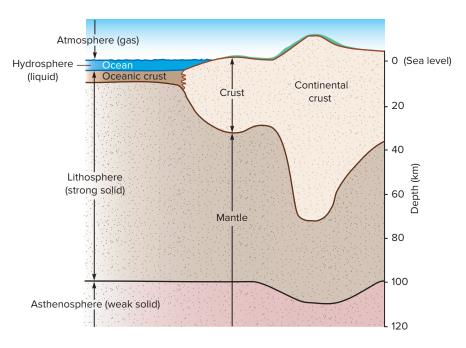


Figure 2.4 Upper layers of Earth may be recognized (1) compositionally, as lower-density crust separated from the underlying higher-density mantle, or (2) on the basis of strength, as rigid lithosphere riding atop "soft plastic" asthenosphere. Notice that the lithosphere includes both the crust and the uppermost mantle.

The differences in strength and mechanical behavior between solid, "plastic," and fluid states are partly responsible for earthquakes and volcanoes.

BEHAVIOR OF MATERIALS

abb91168 ch02 023-046.indd 27

The concepts of gas, liquid, and solid are familiar. Gases and liquids are both **fluids**, but a gas is capable of indefinite expansion, while a liquid is a substance that flows readily and has a definite volume but no definite shape. A solid is firm; it offers resistance to pressure and does not easily

change shape. What is not stated but is implicit in these definitions is the effect of time. All of these definitions describe behavior at an instant in time—but how do the substances behave when viewed over a longer timescale? Specifically, some solids yield to long-term pressure such that at any given moment, they are solid, yet internally they are deforming and flowing—that is, behaving as a fluid. A familiar example is the ice in a glacier. When a glacier is hit with a rock hammer, solid chunks of brittle ice break off. Yet, inside the glacier, atoms are changing positions within the ice and dominantly moving to downhill positions of lower



The Layered Earth

27



Side Note

Volcanoes and the Origin of the Ocean, Atmosphere, and Life

The **elements** in volcanic gases are predominantly hydrogen (H), oxygen (O), carbon (C), sulfur (S), chlorine (Cl), and nitrogen (N). These gaseous elements combine at Earth's surface to make water (H_2O), carbon dioxide (CO_2), sulfur dioxide (SO_2), hydrogen sulfide (H_2) with its rotten egg smell, carbon monoxide (CO), nitrogen (IV_2), hydrogen (IV_2), hydrochloric acid (IV_2), methane (IV_2), and numerous other gases. The dominant volcanic gas is water vapor; it commonly makes up more than 90% of total gases.

The elements of volcanic gases (C, H, O, N, S, Cl) differ from the elements of volcanic rocks: oxygen (O), silicon (Si), aluminum (Al),

iron (Fe), calcium (Ca), magnesium (Mg), sodium (Na), and potassium (K). The elements of volcanic gases make up the oceans, the atmosphere, and life on Earth, but they are rare in rocks. The 4.5 billion years of heat flow from Earth's interior have "sweated" out many lightweight elements and brought them to the surface via volcanism. Billions of years of volcanism on Earth go a long way toward explaining the origin of the continents, the oceans, the present atmosphere, and the surface concentration of the CHON elements (carbon, hydrogen, oxygen, nitrogen) of which all life on Earth is composed and on which it depends.

stress. At no instant in time does the glacier fit our everyday concept of a liquid, yet over time, the glacier is flowing downhill as an ultra-high-viscosity fluid.

When materials are subjected to sufficient **stress**, or force, they deform or undergo **strain** in different ways (figure 2.5). Stress may produce **elastic** (or recoverable) deformation, as when you pull on a spring. The spring deforms while you pull or stress it, but when you let go, it recovers and returns to its original shape.

If greater stress is applied for a longer time or at higher temperatures, **ductile** (or plastic) deformation may occur, and the change is permanent. You can visualize this with a wad of chewing gum or Silly Putty. If you squeeze them in your hands, they deform. Set them down and they stay in the deformed shape; this is ductile deformation. Another example occurs deep within glaciers where the ice deforms and moves with ductile flow.

If stress is applied rapidly to a material, it may abruptly fracture or break into pieces; this is **brittle** deformation. Take a chunk of ice from your refrigerator and drop it or hit it; it will shatter with brittle failure. Notice that the ice in a glacier exhibits both brittle and ductile behavior. Near the surface, there is little pressure on the rigid ice and it abruptly fractures when stressed. Deep within the glacier, where the weight of overlying ice creates a lot of pressure, the ice deforms and moves by ductile flow. The style of ice behavior depends on the amount of pressure confining it.

The type of mechanical behavior illustrated by ice deep within a glacier typifies that of the rock within the Earth's mantle. This rock is **plastic** in the sense used by William James in his 1890 *Principles of Psychology*. He defined *plastic* as "possession of a structure weak enough to yield to an influence, but strong enough not to yield all at once."

When a material such as rock is subjected to the same large amounts of stress on all sides, it compresses. When stresses coming from different directions vary, strain can occur. When the differences in stress are low, strain is elastic and reversible. As stress differences increase, the **yield stress** is reached and permanent strain occurs. Most rocks

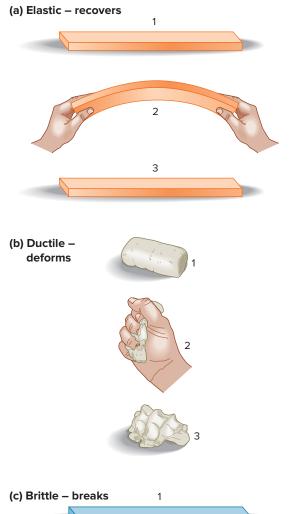
are brittle at the low temperatures and low pressures at Earth's surface. Most rocks are ductile at the high temperatures and high pressures at depth inside Earth. In the asthenosphere, rock deforms in a "soft plastic" fashion. Most of the deeper mantle rock is solid but not brittle; it behaves as a "stiff plastic"—it deforms.

The top of the asthenosphere comes to the surface at some of the ocean's volcanic mountain chains but lies more than 100 km (about 60 mi) below the surface in other areas. It has gradational upper and lower boundaries and is about 250 km (155 mi) thick. What are the effects of having this "soft plastic" ductile zone so near Earth's exterior? Within the asthenosphere, there is a lot of flowage of rock that helps cause Earth's surface to rise and fall. For example, Earth is commonly described as a sphere, but it is not. Earth may be more properly described as an oblate ellipsoid that is flattened at the poles (nearly 30 km, or 19 mi) and bulged at the equator (nearly 15 km, or 9 mi). Earth is neither solid enough nor even strong enough to spin and maintain a spherical shape. Rather, Earth deforms its shape in response to the spin force. The flattening of Earth during rotation is analogous to the flattening of the early Solar System from a sphere to a disk (see figure 2.1).

ISOSTASY

The Earth is not a homogeneous, solid ball. It is instead a series of floating layers where less dense materials successively rest upon layers of more dense materials (see figure 2.3). The core, with densities up to 16 g/cm³, supports the mantle, with densities ranging from 5.7 to 3.3 g/cm³. Atop the denser mantle float the continents, with densities around 2.7 g/cm³, which in turn support the salty oceans, with densities of about 1.03 g/cm³, and then the least dense layer of them all—the atmosphere. The concept of floating layers holds true on smaller scales as well. For example, the oceans are made of layered masses of water of differing densities. Very cold, dense Antarctic waters flow along the ocean bottoms and are overlain by cold Arctic water, which





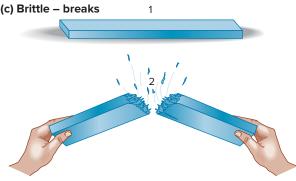
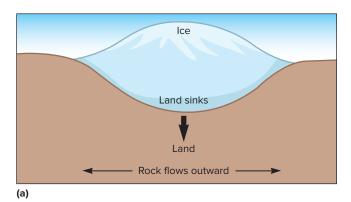


Figure 2.5 Behavior of materials. (a) Elastic: bend a thin board; let it go and the board recovers its original shape. (b) Ductile: squeeze a wad of bubblegum or Silly Putty; let it go and the mass stays in the deformed shape. (c) Brittle: bend a thin board sharply and it breaks.

is overlain by extra-salty waters, which in turn are overlain by warmer, less dense seawater. Earth is composed basically, from core through atmosphere, of density-stratified layers.

The concept of **isostasy** was developed in the 19th century. It applies a principle of **buoyancy** to explain how the low-density continents and mountain ranges literally float

on the denser mantle below. Just as an iceberg juts up out of the ocean while most of its floating mass is beneath sea level, so does a floating continent jut upward at the same time it has a thick "root" beneath it (see figure 2.4). Visualize a boat floating in water: Add a load onto the boat and it sinks downward; remove the load and the boat rises upward. So it is with a continent. Add a load onto the land, such as a large glacial ice mass, and the land will sink downward as rock at depth flows outward within the asthenosphere; remove the load (the ice melts), and the land rises or rebounds upward as rock flows inward in the asthenosphere (figure 2.6). An example of this buoyancy effect, or isostatic equilibrium, was defined by carefully surveying the landscape before and after the construction of Hoover Dam across the Colorado River east of Las Vegas, Nevada. On 1 February 1935, the impoundment of Lake Mead began. By 1941, about 24 million acre feet of water had been detained, placing a weight of 40,000 million tons over an area of 232 square miles. Although this is an impressive reservoir on a human scale, what effect can you imagine it having on the whole Earth? In fact, during the 15 years from 1935 to 1950, the central region beneath the reservoir sank up to 175 mm (7 in) (figure 2.7). The relatively simple act of impounding water behind the dam triggered an isostatic adjustment as asthenosphere rock flowed away from the pressure of the overlying reservoir, causing the area to subside.



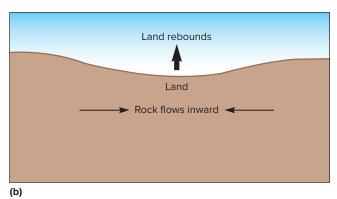


Figure 2.6 Isostatic equilibrium. (a) Land sinks as weight of ice causes rock at depth to flow outward. (b) Land rebounds as ice melts and removes weight, causing rock at depth to flow inward.

The Layered Earth



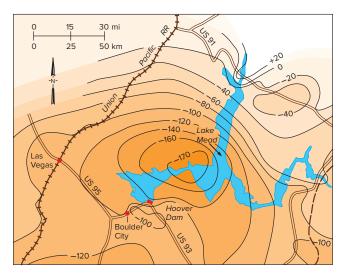


Figure 2.7 Isostatic downwarping caused by the weight of the water in Lake Mead, from 1935 to 1950. Black circular lines (contour lines) define the depressed land surface. In the center is a –170 line where land sank 170 mm (7 in).

Source: Smith, W.O. et al. Comprehensive Survey of Sedimentation in Lake Mead, 1948-49, in US Geological Survey Professional Paper 295, 1960.

Just how solid and firm is the surface of the Earth we live on? Larger-scale examples are provided by the great ice sheets of the recent geologic past. The ice sheet that buried the Finland-Sweden region was up to 3 km (2 mi) thick less than 20,000 years ago. The land was depressed beneath this great weight. By 10,000 years ago, the ice sheet had retreated and melted, and the water returned to the ocean. The long-depressed landmass, now freed from its heavy load, is rebounding upward via isostatic adjustment. In the past 10,000 years, northeastern coastal Sweden and western Finland have risen about 200 m (650 ft). This upward movement was vividly shown during excavation for a building foundation in Stockholm, Sweden. Workers uncovered a Viking ship that had sunk in the harbor and been buried with mud. The ship had been lifted above sea level, encased in its mud shroud, as the harbor area rose during the ongoing isostatic rebound. Gravity measurements of this region show a negative anomaly, indicating that another 200 m (650 ft) of isostatic uplift is yet to come. The uplift will add to the land of Sweden and Finland and reduce the size of the Gulf of Bothnia between them.

Some of the early uplifting of land after ice-sheet removal occurred in rapid movements that ruptured the ground surface, generating powerful earthquakes. In northern Sweden, there are ground ruptures up to 160 km (100 mi) long with parallel cliffs up to 15 m (50 ft) high. The rocks in the region are ancient and rigid, suggesting that ruptures may go 40 km (25 mi) deep and that they generated truly large earthquakes.

Vertical movements of the rigid lithosphere floating on the flexible asthenosphere are well documented. If we add a load on the surface of Earth, we can measure the downward movement. For example, a 100 meter thick ice mass will cause the land to sink about 27.5 m (90 ft). Antarctica is buried beneath ice up to 4,470 m (2.8 mi) thick. Thus, Antarctica is depressed up to 1,230 m (4,000 ft), placing most of the continent below sea level. If the ice is removed, Antarctica will slowly rise up and become the fifth largest landmass on Earth. The surface of Earth clearly is in a delicate vertical balance. Do major adjustments and movements also occur horizontally? Yes, there are horizontal movements between the lithosphere and the asthenosphere, which will bring us into the realm of plate tectonics (described later in this chapter).

Internal Sources of Energy

The flow of energy from Earth's interior to its surface comes mainly from three sources: impact energy, gravitational energy, and the ongoing decay of radioactive isotopes.

IMPACT ENERGY AND GRAVITATIONAL ENERGY

The impact energy of masses colliding with the growing Earth produced heat. Tremendous numbers of large and small asteroids, meteorites, and comets hit the early Earth, their energy of motion being converted to heat on impact.

Gravitational energy was released as Earth pulled into an increasingly dense mass during its first tens of millions of years. The ever-deeper burial of material within the growing mass of Earth caused an increasingly greater gravitational pull that further compacted the interior. This gravitational energy was converted to heat.

The immense amount of heat generated during the formation of Earth did not readily escape because rock conducts heat very slowly. Some of this early heat is still flowing to the surface today.

RADIOACTIVE ISOTOPES

Energy is released from **radioactive isotopes** as they decay. Radioactive isotopes are unstable and must kick out subatomic particles to attain stability. As radioactive isotopes decay, heat is released.

In the beginning of Earth, there were abundant, short-lived radioactive isotopes, such as aluminum-26, that are now effectively extinct, as well as long-lived radioactive isotopes, many of which have now expended much of their energy (table 2.1). Young Earth had a much larger complement of radioactive isotopes and a much greater heat production from them than it does now (figure 2.8). With a declining output of radioactive heat inside the Earth, the flow of energy from Earth's interior is on a slow decline heading toward zero.



TABLE 2.1			
Some Radioactive Isotopes in Earth			
Parent	Decay Product	Half-Life (billion years)	
Aluminum-26	Magnesium-26	0.00072 (720,000 years)	
Uranium-235	Lead-207	0.71	
Potassium-40	Argon-40	1.3	
Uranium-238	Lead-206	4.5	
Thorium-232	Lead-208	14	
Rubidium-87	Strontium-87	48.8	
Samarium-147	Neodymium-143	106	

The radioactive-decay process is measured by the **half-life**, which is the length of time needed for half the present number of atoms of a radioactive isotope (parent) to disintegrate to a decay product. As the curve in figure 2.9 shows, during the first half-life, one-half of the atoms of the radioactive isotopes decay. During the second half-life, one-half of the remaining radioactive atoms decay (equivalent to 25% of the original parent atoms). The third half-life witnesses the third halving of radioactive atoms present (12.5% of the original parent atom population), and so forth. Half-lives plotted against time produce a negative exponential curve; this is the opposite direction of a positive exponential curve, such as interest being paid on money in a savings account.

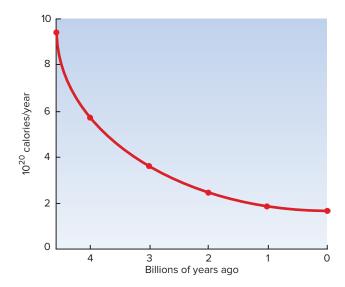


Figure 2.8 The rate of heat production from decay of radioactive atoms has declined throughout the history of Earth.

The sum of the internal energy from impacts, gravity, and radioactive isotopes, plus additional energy produced by **tidal friction**, is very large. The greater abundance of radioactive isotopes at Earth's beginning, along with the early gravitational compaction and more frequent meteorite impacts, combined to elevate Earth's internal temperature during its early history. It is noteworthy that this heat buildup reached a maximum early in Earth's history and has declined significantly since then. Nonetheless, the flow of internal heat toward Earth's surface today is still great enough to provide the energy for continents to drift, volcanoes to erupt, and earthquakes to shake.

Internal Sources of Energy

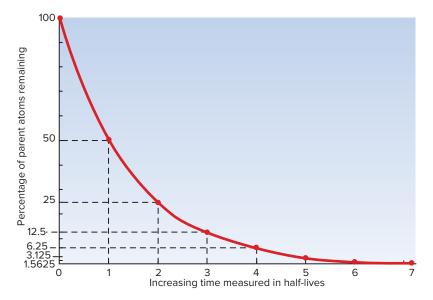


Figure 2.9 Negative exponential curve showing decay of radioactive parent atoms to stable daughter atoms over time. Each half-life witnesses the disintegration of half the remaining radioactive parent atoms.



In Greater Depth

Radioactive Isotopes

Each chemical element has a unique number of positively charged protons that define it. However, the number of neutrons varies, giving rise to different forms of the same element, known as isotopes. Some isotopes are radioactive and release energy during their decay processes. In radioactive decay, unstable parent atoms shed excess subatomic particles, reducing their weight and becoming smaller atoms (figure 2.10). The overly heavy radioactive isotopes slim down to a stable weight by splitting apart, as in emitting alpha particles consisting of two protons and two neutrons (effectively, the nucleus of a helium atom). Beta particles are electrons freed upon a neutron's splitting. Gamma radiation, which is similar to X-rays but with shorter wavelength, is emitted, lowering the energy level of a nucleus. As the rapidly expelled particles are slowed and absorbed by surrounding matter, their energy of motion is transformed into heat.

DATING THE EVENTS OF HISTORY

The same decaying radioactive isotopes producing heat inside Earth, Moon, and meteorites also may be read as clocks that date events in history. For example, uranium-238 decays to lead-206 through numerous steps involving different isotopes and new elements (figure 2.11). By emitting alpha and beta particles, 32 of the 238 subatomic particles in the U-238 nucleus are lost, leaving behind the 206 particles of the Pb-206 nucleus. Laboratory measurements of the rate of the decay process have given us the U-238-to-Pb-206 half-life of 4.5 billion years. These facts may be applied to quantifying history by reading the radiometric clocks preserved in some minerals. For example, some igneous rocks (crystallized from magma) can be crushed, and the very hard mineral zircon can be separated from them. Zirconium, the diamond substitute in jewelry, is synthesized from zircon. Zircon crystals contain uranium-238 that was locked into their atomic structure when they crystallized from magma, but they originally contained virtually no lead-206. Thus, the lead-206 present in the crystal must have come from decay of uranium-238.

The collected zircon crystals are crushed into a powder and dissolved with acid under ultraclean conditions. The sample is placed in a mass spectrometer to measure the amounts of parent uranium-238 and lead-206 present. Then, with three known values—(1) the amount of U-238, (2) the amount of Pb-206, and (3) the half-life of 4.5 billion years for the decay process—it is easy to calculate how long the U-238 has been decaying into Pb-206 within the zircon crystal. In other words, the calculation tells us how long ago the zircon crystal formed and consequently the time of formation of the igneous rock.

See the Geologic Timescale in Chapter 18 and Appendix A.

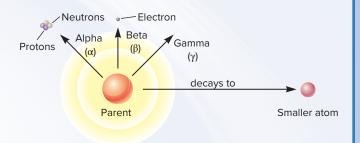


Figure 2.10 A radioactive parent atom decays to a smaller atom by emitting alpha particles (such as the nucleus of a helium atom, i.e., two protons and two neutrons), beta particles (electrons), and gamma radiation (such as X-rays).

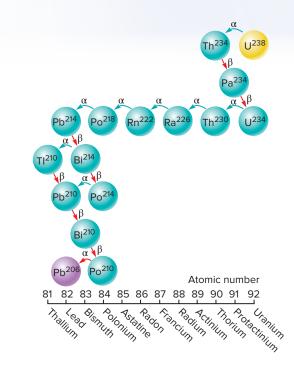


Figure 2.11 Radioactive uranium-238 (U²³⁸) decays to stable lead-206 (Pb²⁰⁶) by steps involving many intermediate radioactive atoms. The atomic number is the number of protons (positively charged particles) in the nucleus.

Age of Earth

The oldest Solar System materials are about 4.57 billion (4,570 million) years old. The 4.57-billion-year age has been measured using radioactive isotopes and their decay products collected from meteorites. In 2016, analyses of Moon rocks for hafnium (¹⁸²HF) indicated that the Moon had formed by

4.51 billion years ago. The oldest ages obtained on Earth materials are 4.4 billion years, measured on sand grains of the mineral zircon collected from within a 3.1-billion-year-old sandstone in western Australia. The oldest Earth rocks found to date are in northwest Canada; they are 4.055 billion years old. These rocks are of crustal composition, implying that they were recycled and formed from even older rocks.



In Greater Depth

Radioactivity Disasters

The term radioactivity disasters brings to mind the meltdown of the uranium-rich core of a nuclear-power plant, as happened at Chernobyl in Ukraine, part of the former Soviet Union, on 26 April 1986. This human-caused disaster occurred when the night-shift workers made a series of mistakes that unleashed a power surge so great that the resultant explosions knocked off the 1,000-ton lid atop the nuclear reactor core, blew out the building's side and roof, triggered a partial meltdown of the reactor core's radioactive fuel, and expelled several tons of uranium dioxide fuel and fission products, including cesium-137 and iodine-131, in a 5 km (3 mi) high plume. As many as 185 million curies of radioactivity were released. (The worst U.S. incident released 17 curies from the Three Mile Island nuclear-power plant in Pennsylvania in 1979.) After the 1:24 a.m. explosion, people near Chernobyl were at least fortunate that they were indoors and thus somewhat sheltered, there was no rain in the area, and the contaminant plume rose high instead of hugging the ground. The cloud of radioactive contaminants affected people, livestock, and agriculture from Scandinavia to Greece. In the Chernobyl power-plant area, about 50 people died directly. Most of the deaths would come later from cancer and other diseases. The worst contaminant is radioactive iodine-131, which lodges in the thyroid. Cancer of the thyroid is expected to be common in the area; it is estimated that it will shorten the lives of about 8,000 people.

An earthquake may have helped trigger this disaster. The Chernobyl power-plant workers were having difficulties in the early morning hours of 26 April, and then a magnitude 3 earthquake occurred 12 km (7 mi) away. The panicked supervisor thought the shaking meant the power plant was losing control, and he quickly implemented emergency maneuvers, but they jammed the internal works of the reactor, leading to the fateful explosion 22 seconds after the earthquake.

Earthquakes and radioactivity disasters entered the news again on 11 March 2011 when a great magnitude 9 earthquake and resultant tsunami in Japan destroyed several nuclear-power plants. This event is described in chapter 8 on tsunami.

Chernobyl was a human-caused disaster. What can happen under natural conditions? Today, on Earth and Moon, uranium is present mostly as the heavier U-238 isotope, which has a combined total of 238 protons and neutrons in each uranium atom nucleus. The lighter-weight uranium isotope, U-235, makes up only 0.7202% of all uranium atoms. In nuclear-power plants, the uranium ore fed to nuclear reactors is enriched to 2–4% U-235 to promote more potent reactions. Remember from table 2.1 that U-235 has a half-life of 0.71 billion years, whereas the half-life of U-238 is 4.5 billion years. Because U-235 decays more rapidly, it would have been relatively more abundant in the geologic past. In fact, at some past time, the U-235 natural percentage relative to U-238 would have been like the U-235 percentage added to U-238 and fed as ore to nuclear reactors today.

Have natural nuclear reactors operated in the geologic past? Yes. A well-documented example has been exposed in the Oklo uranium mine near Franceville in southeastern Gabon, a coastal country in equatorial West Africa. At Oklo, 2.1 billion years ago, sands and muds accumulated along with organic carbon from the remains of fossil bacteria. These carbon-bearing **sediments** were enriched in uranium; U-235 was then 3.16% of total uranium. The sand and mud sediments were buried to shallow depths, and at least 800 m³ (1,050 yd³) of uranium ore sustained **nuclear fission** reactions that generated temperatures of about 400°C (750°F) regionally and much higher temperatures locally. At Oklo, 17 sites started up as natural nuclear reactors about 1.85 billion years ago; they ran for at least 500,000 years (and maybe as long as 2 million years). Nine of the natural reactors that have been carefully studied are estimated to have produced at least 17,800 megawatt years of energy.

Our understanding of the age of Earth is improving rapidly as new technologies allow measurement of more types of radioactive isotopes. It now seems that Earth has existed as a coherent mass for about 4.54 billion years. Earth must be younger than the 4.57-billion-year-old materials that collided and clumped together to form the planet. The time it took to build Earth is possibly as short as 30 million years. The collision of Earth with the Mars-size body that led to forming our Moon seems to have occurred between 4.537 and 4.533 billion years ago, suggesting that Earth was already a large, coherent mass at that time. Coming from the other direction, Earth must be older than the 4.4-billion-year-old zircon grains collected from sand-stone in Australia. In sum, our planet has existed for about 4.5 billion years.

The work to exactly determine the age and early history of Earth continues today. It is challenging to try and find the oldest minerals and rocks because Earth is such an energetic planet that surface rocks are continually being formed and destroyed. Because of these active Earth processes, truly old materials are rarely preserved; there have been too many events over too many years that destroy rocks.

Plate Tectonics

The grand recycling of the upper few hundred kilometers of Earth is called the **tectonic cycle.** The Greek word *tekton* comes from architecture and means "to build"; it has been adapted by geologists as the term **tectonics**, which describes the building of **topography** and the deformation and movement within Earth's outer layers.

Adding the horizontal components of movements on Earth allows us to understand the tectonic cycle. Ignoring complexities for the moment, the tectonic cycle can be simplified as follows (figure 2.12). First, the asthenosphere rises and melts to form magma that flows upward and cools to form new ocean floor/lithosphere. Second, the new

Plate Tectonics



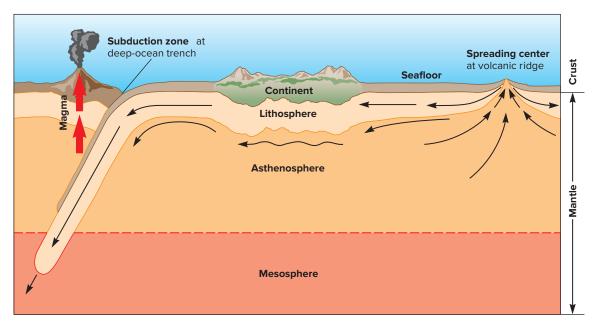


Figure 2.12 Schematic cross-section of the tectonic cycle. Magma rises from the asthenosphere to the surface at the oceanic volcanic ridges where it solidifies and adds to the plate edges. As the igneous rock cools, the plate subsides and gravity pulls the plates from their topographic highs. The plate continues to cool, grows thicker at its base, becomes denser, collides with a less-dense plate, and turns down into the mantle, where it is ultimately reassimilated.

lithosphere slowly moves laterally away from the zones of oceanic crust formation on top of the underlying asthenosphere; this phenomenon is known as **seafloor spreading.** Third, when the leading edge of a moving slab of oceanic lithosphere collides with another slab, the older, colder, denser slab turns downward and is pulled by gravity back into the asthenosphere, a process called **subduction**, while the less-dense, more buoyant slab overrides it. Last, the slab pulled into the mantle is reabsorbed. The time needed to complete this cycle is long, commonly in excess of 250 million years.

If we adopt the perspective of a geologist-astronaut in space and look down upon the tectonic cycle, we see that the lithosphere of Earth is broken into pieces called **plates** (figure 2.13). The study of the movements and interactions of the plates is known as **plate tectonics**. The gigantic pieces of lithosphere (plates) pull apart during seafloor spreading at **divergence zones**, slide past at **transform faults**, or collide at **convergence zones**. At divergence zones, plates are constructed. At transform faults, plates are conserved. At convergence zones, plates are destroyed. These plate-edge interactions are directly responsible for most of the earth-quakes, volcanic eruptions, and mountains on Earth.

Another way that plate tectonics can be visualized is by using a hard-boiled egg as a metaphor for Earth. Consider the hard-boiled egg with its brittle shell as the lithosphere, the slippery inner lining of the shell as the asthenosphere, the egg white (albumen) as the rest of the mantle, and the yolk as the core. Before eating a hard-boiled egg, we break its brittle shell into pieces that slip around as we try to pluck them off. This hand-held model of brittle pieces

being moved atop a softer layer below is a small-scale analogue to the interactions between Earth's lithosphere and asthenosphere.

DEVELOPMENT OF THE PLATE TECTONICS CONCEPT

Our planet is so large and so old that the combined efforts of many geologists and philosophers over the past few hundred years have been required to amass enough observations to begin understanding how and why Earth changes as it does. The first glimpse of our modern understanding began after the European explorers of the late 1400s and 1500s made maps of the shapes and locations of the known continents and oceans. These early world maps raised intriguing possibilities. For example, in 1620, Francis Bacon of England noted the parallelism of the Atlantic coastlines of South America and Africa and suggested that these continents had once been joined. During the late 1800s, the Austrian geologist Eduard Suess presented abundant evidence in support of Gondwanaland, an ancient southern supercontinent composed of a united South America, Africa, Antarctica, Australia, India, and New Zealand, which later split apart. This process of the continents moving, splitting, and recombining is known as continental drift. The most famous and outspoken of the early proponents of continental drift was the German meteorologist Alfred Wegener. In his 1915 book, The Origin of Continents and Oceans, he collected all available evidence, such as similar rocks, fossils, and geologic structures, on opposite sides of the Atlantic Ocean. Wegener suggested that all the continents had once been united in a



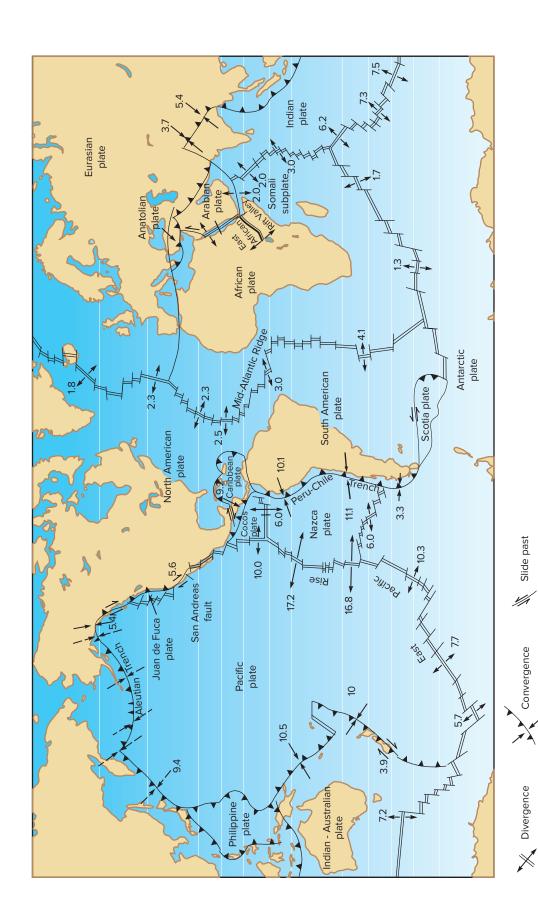


Figure 2.13 Map of the major tectonic plates with arrows showing directions of movement. Rates of movement are shown in centimeters per year.

Plate Tectonics

35

abb91168_ch02_023-046.indd 35 03/11/21 10:13 PM



In Greater Depth

Earth's Magnetic Field

Anyone who has ever held a compass and watched the free-turning needle point toward the north has experienced the **magnetic field** that surrounds Earth. The Chinese invented and were the first to use magnetic compasses. They in turn taught 14th-century European travelers, who brought this knowledge back to Europe, where it was developed into the navigational tool that helped late 15th-century explorers make their voyages of discovery.

Earth's magnetic field operates as if a gigantic bar magnet were located in the core of Earth inclined 11° from vertical (figure 2.14). The **magnetic pole** and geographic North Pole do not coincide, but the magnetic pole axis has apparently always been near the rotational pole axis. Notice in figure 2.14 that the inclination of the magnetic lines of force with respect to Earth's surface varies with **latitude**. At the magnetic equator, the magnetic lines of force are parallel to Earth's surface (inclination of 0°). Toward the poles, either northward or southward, the angle of inclination continuously increases until it is perpendicular to the surface at both the north and south magnetic poles (inclinations of 90°). Notice also that the lines of force are inclined downward and into Earth's surface near the North Pole and upward and out of Earth's surface near the South Pole.

In reality, the interior of Earth is much too hot for a bar magnet to exist. **Magnetism** in rocks is destroyed by temperatures above 550°C (1,020°F), and temperatures in Earth's core are estimated to reach 5,800°C (10,470°F). The origin of Earth's magnetic field involves movements of the iron-rich fluid in the outer core, which generate electric currents that in turn create the magnetic field. Fluid iron is an excellent conductor of electricity. The molten iron flowing around the solid inner core is a self-perpetuating dynamo deriving its energy both from the rotation of Earth and from the convection of heat released by the crystallization of minerals at the boundary of the inner and outer cores.

A closer look at Earth's magnetic field yields several problems awaiting resolution. The simplified magnetic field portrayed in figure 2.14 does not show the complexities that occur over years and centuries as the magnetic field's strength waxes and wanes. More than 400 years of measurements document variations in the strength and stability of the magnetic field. At present, the strength is 10% weaker than in the year 1845, but the field strength is still about twice as strong as the long-term average. The flow of fluid iron in the outer core has regions of turbulence, including motions as complex as whirlpools. Change is normal, and in turn, the magnetic field fluctuates.

In addition, the magnetic pole moves about the geographic North Pole region in an irregular pattern. The rapid rotation of Earth holds the magnetic pole near the pole of rotation, but the magnetic pole wanders enough that it crosses 5° to 10° of latitude each century. In recent decades, the magnetic pole has moved at rates of 10 to 40 km (6 to 25 mi) per year, with intervals up to 60 km (37 mi) per year.

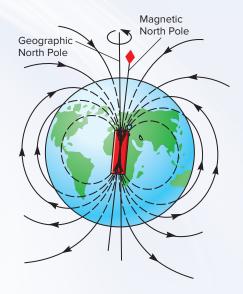


Figure 2.14 Schematic diagram of Earth's magnetic field. The bar magnet pictured does not exist, but it would create the same magnetic field achieved by the electric currents in Earth's liquid, iron-rich outer core. Notice that (1) the magnetic pole and the rotational pole do not coincide, (2) the magnetic lines of force are parallel to Earth's surface at the magnetic equator and perpendicular at the magnetic poles, and (3) the lines of force go into Earth at the North Pole and out at the South Pole.

Every several thousand to tens of millions of years, a highly dramatic change occurs in the magnetic field: the magnetic polarity reverses. In a reversal, the orientation of the magnetic field flipflops from a north (normal) polarity to a south (reverse) polarity or vice versa. It has been 780,000 years since the last long-term reversal. Models run on supercomputers indicate that reversals take less than a thousand years to complete. During a reversal, it appears that the magnetic field does not disappear; it just gets more complex. The magnetic lines of force become twisted and tangled, but a magnetic field still exists to protect life on Earth from much of the incoming solar and space radiation.

The change in orientation of the magnetic field leaves its imprint in rocks, where geologists (paleomagnetists) can read it. The paleomagnetic history contained in the rocks has provided the most important evidence of seafloor spreading; it also has allowed charting of the paths of continents as they have moved through different latitudes. In addition, the record of magnetic reversals provides the data for a magnetic timescale, a third geologic timescale. (The first timescale is based on the irreversible sequence of fossils occurring in sedimentary rocks, and the second timescale is founded on the decay of radioactive isotopes.)

supercontinent called **Pangaea** (pan meaning "all" and gaea meaning "earth").

Much is made of the fact that during his lifetime, Wegener's hypothesis of continental drift garnered more ridicule than acceptance. But why were his ideas not widely accepted? Wegener presented an intriguing hypothesis well supported with observations and logic, but his mechanism was deemed impossible. Geologists and geophysicists could not visualize how a continent could break loose from the underlying rocks and plow a path over them. The breakthrough in understanding came when the ocean floors were studied and the data were best explained by the formation



of new seafloor that spread apart and later was consumed by subduction. When scientists realized that the lithosphere decouples from the asthenosphere and moves laterally, they understood how the relatively small, low-density continents, set within the oceanic crust, could be carried along as incidental passengers (see figure 2.12).

In the mid-1960s, evidence abounded, mechanisms seemed plausible, and the plate tectonic theory was developed and widely accepted. Wegener was restored to an elevated status. Scientific understanding grew with the addition of new data, old hypotheses were modified, and new theories were created. Science is never static; it is a growing, evolving body of knowledge that creates ever-better understanding of how Earth works.

It is rare in science to find widespread agreement on a large-scale hypothesis such as plate tectonics. But when data from Earth's magnetic field locked inside seafloor rocks were widely understood, skeptics around the world became convinced that seafloor spreading occurs and that the concept of plate tectonics is valid. These paleomagnetic data are so powerful that we need to understand their story so that plate tectonics can be seen as real.

MAGNETIZATION OF VOLCANIC ROCKS

Lava is magma that erupts from a volcano, flows outward as a sheetlike mass, slows down, and stops. Then, as the lava cools, minerals begin to grow as crystals. Some of the earliest formed crystals incorporate iron into their structures. After the lava cools below the Curie point, about 550°C (1,020°F), atoms in iron-bearing minerals become magnetized in the direction of Earth's magnetic field at that time and place. The lined-up atoms in the iron-rich crystals behave like compass needles pointing toward the magnetic pole of their time (measured as declination or "compass bearing"); they also become inclined at the same angle as the lines of force of the magnetic field (measured as inclination or dip). Ancient magnetic fields have been measured in rocks as old as 3.5 billion years.

Lava flows pile up as sequences of stratified (layered) rock, and the magnetic polarity of each rock layer can be measured (figure 2.15). Many of the volcanic rocks also contain minerals with radioactive isotopes that allow us to determine the age of the volcanic rock—that is, how long ago the lava flow solidified. When this information is plotted together in a vertical column, a timescale of magnetic polarities emerges (figure 2.16). It is interesting to note that

Plate Tectonics



Figure 2.15 A stratified pile of former lava flows of the Columbia River Basalt exposed in the east wall of Grand Coulee, Washington. The oldest flow is on the bottom and is overlain by progressively younger flows.

©University of Washington Libraries, Special Collections, John Shelton Collection, Shelton KC10288



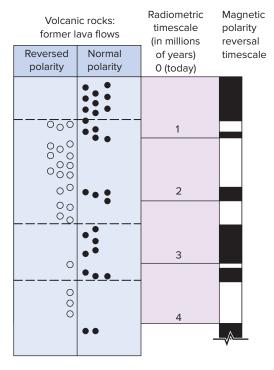


Figure 2.16 A portion of the magnetic polarity timescale. Magnetic polarity measurements in volcanic rocks combined with radiometric ages determined from the same rocks allow formation of a timescale based on magnetic polarity reversals. Notice the unique and nonrepetitive pattern of the polarity reversals.

Source: From P. J. Wyllie, The Way the Earth Works. Copyright @ 1976 John Wiley & Sons, Inc., New York.

the timing of polarity reversals appears to be random. There is no discernible pattern to the lengths of time the magnetic field was oriented either to the north or to the south. The processes that reverse the polarity of the magnetic field are likely related to changes in the flow of the iron-rich liquid in the outer core. The reversal-causing mechanism does not occur at any mathematically definable time interval.

Magnetization Patterns on the Seafloors

Since the late 1940s, oceanographic research vessels crisscrossing the Atlantic Ocean have towed magnetometers to measure the magnetization of the seafloor. As the number of voyages grew and more data were obtained, a striking pattern began to emerge (figure 2.17). The floor of the Atlantic Ocean is striped by parallel bands of magnetized rock that show alternating polarities. The pattern is symmetrical and parallel with the midocean volcanic **ridge** (spreading center). That is, each striped piece of seafloor has its twin on the other side of the oceanic mountain range.

A remarkable relationship exists between the time of reversals of magnetic polarity, as dated radiometrically from a sequence of solidified lava flows (see figure 2.16), and the widths of alternately polarized seafloor (figure 2.17)—they are comparable. How stunning it is that the widths of magnetized seafloor strips have the same ratios as the lengths of time between successive reversals of Earth's magnetic field. This means that distance in kilometers is proportional to time in millions of years. Now, if Earth's magnetic field

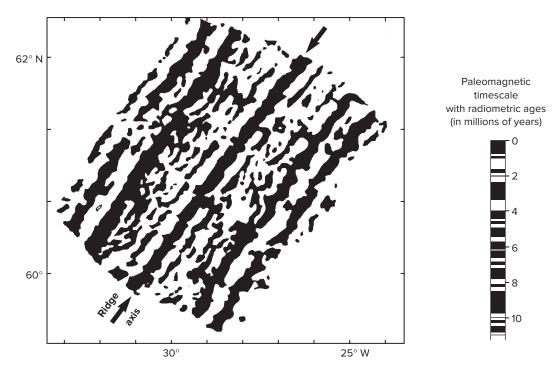


Figure 2.17 Map of the magnetically striped Atlantic Ocean floor southwest of Iceland. Black areas are magnetized pointing to a north pole and white areas to a south pole. Notice the near mirror images of the patterns on each side of the volcanic ridge (spreading center).



is reversing polarity in a known timescale and if that timescale reappears in distances, then the relationship must take the form of a velocity. That is, magma is injected into the oceanic ridges where it is imprinted by Earth's magnetic field as it cools to form new rock. Then the seafloor/ocean crust/lithosphere is physically pulled away from the oceanic ridges as if they were parts of two large conveyor belts going in opposite directions (figure 2.18).

The evidence provided by the paleomagnetic timescale and the magnetically striped seafloors is compelling. These phenomena are convincing evidence that seafloor spreading occurs and that plate tectonics is valid.

Let's now consider other evidence supporting plate tectonics.

Earthquake Evidence

The map of earthquake **epicenters** (figure 2.19) can be viewed as a connect-the-dots puzzle. Each epicenter represents a place where one major section of rock has moved past another section. Take your pen or pencil,

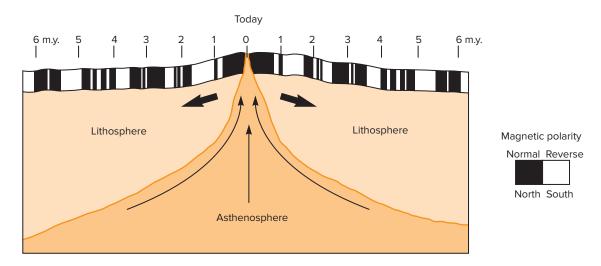


Figure 2.18 Cross-section of magnetically striped seafloor. Numbers above the seafloor are radiometrically determined ages in millions of years (m.y.). The near mirror-image magnetic pattern is like a tape recorder that documents "conveyor belt" movements away from volcanic ridges.

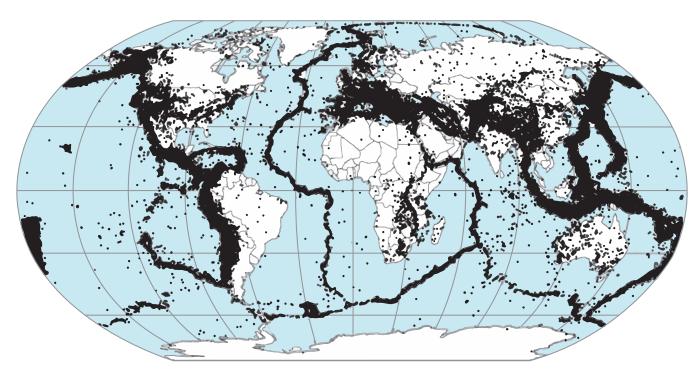


Figure 2.19 Map of earthquake epicenters, 1963–1998. Notice that epicenters are concentrated in linear belts. Connecting the "dots" defines the tectonic plates.

Source: NASA, DTAM project team

Plate Tectonics